

An OSSE-based Evaluation of Hybrid Variational-Ensemble Data Assimilation for the NCEP GFS, Part II: 4DEnVar and Hybrid Variants

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Revised 12 October 2014

Monthly Weather Review

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Abstract

1 This work describes the formulation of a hybrid 4DEnVar algorithm and initialization
2 options utilized within the National Centers for Environmental Prediction global data
3 assimilation system. Initialization schemes that are proposed for use are the tangent linear
4 normal mode constraint, weak constraint digital filter, and a combination thereof.

5 An observing system simulation experiment is carried out to evaluate the impact of
6 utilizing hybrid 4DEnVar with various initialization techniques. The experiments utilize a dual
7 resolution configuration, where the ensemble is run at roughly half the resolution of the
8 deterministic component. It is found that by going from 3D to 4D, analysis error is reduced for
9 most variables and levels. The inclusion of a time-invariant static covariance when used without
10 a normal mode based strong constraint is found to have a small, positive impact on the analysis.
11 The experiments show that the weak constraint digital filter degrades the quality of analysis, due
12 to the use of hourly states to prescribe high frequency noise. It is found that going from 3D to
13 4D ensemble covariances has relatively larger impact in the extratropics, whereas the original
14 inclusion of ensemble-based covariances was found to have the largest impact in the tropics.
15 The improvements found in going from 3D to 4D covariances in the hybrid EnVar formulation
16 are not as large as was found in Part I from the original introduction of the hybrid algorithm.
17 The analyses generated by the 4D hybrid scheme are found to yield slightly improved
18 extratropical height and wind forecasts, with smaller impacts on other variables and in general in
19 the tropics.

20

21 1. Introduction

22 Four-dimensional variational data assimilation (4DVar) techniques that use tangent linear
23 (Lewis and Derber 1985; Courtier et al. 1994) or linear perturbation models (Rawlins et al. 2007)
24 and their corresponding adjoints have been shown to be powerful natural extensions to the
25 3DVar technique. In fact, 4DVar is the method of choice for initialization of single deterministic
26 numerical weather prediction (NWP) applications at many operational centers (Rabier et al.
27 2000; Rosmond and Xu 2006; Gauthier et al. 2007; Rawlins et al. 2007). One attractive feature
28 of 4DVar is that a dynamic model is used to help impose temporal smoothness and physical
29 constraints on the analysis. Additionally, 4DVar allows for the simultaneous assimilation of
30 asynchronous observations throughout a window at their appropriate times by producing a 4D
31 analysis trajectory (Lorenc and Rawlins 2005). This is in contrast with the 3DVar-FGAT
32 method (Rabier et al. 1998; Lawless 2010), which employs 4D model states at the appropriate
33 time to compute innovations but only solves for a solution at a single time, typically at the center
34 of a window. The major drawbacks to the 4DVar technique are the computational cost,
35 complications related to developing and maintaining linearized forecast models and their
36 corresponding adjoints, and basic assumption of linearity for the incremental formulation which
37 may be particularly problematic for high resolution.

38 Much like 3DVar, 4DVar typically assumes a static error covariance but valid at the
39 beginning of the assimilation window. The tangent linear (TL) and adjoint (AD) models then
40 implicitly evolve this background error covariance as part of the variational solver. This
41 procedure does allow for some flow-dependence, though the quality of the initial static
42 background error covariance can still play a crucial role for short assimilation windows. Further,
43 the use of a dynamic model to constrain the solution does help improve the multivariate aspects.

44 Much like for the 3D case, the development and application of a hybrid En4DVar technique,
45 which utilizes ensemble-based covariances for helping to prescribe the background error
46 covariance at the beginning of the assimilation window, has been shown to be beneficial
47 (Buehner et al. 2010a,b; Zhang and Zhang 2012; Clayton et al. 2013; Kuhl et al. 2013). The
48 drawbacks of such a method are the same as those for 4DVar, namely computational cost and the
49 need for TL and AD models. As in Lorenc (2013), hybrid is used to describe a blended
50 covariance and 4DVar implies the use of TL and AD models.

51 Along the lines of the 4D ensemble-based techniques (Hunt et al. 2004) such as the 4D-
52 Local Ensemble Transform Kalman Filter (LETKF; Hunt et al. 2007), several methods
53 expanding on the idea introduced by Lorenc (2003) have recently been proposed to utilize 4D
54 ensemble perturbations within a variational framework (EnVar, Lorenc 2013) to help solve for a
55 4D-analysis increment without the need for TL and AD models (Liu et al. 2008; Tian et al. 2008;
56 Liu et al. 2009; Tian et al. 2011; Buehner et al. 2010a,b; Liu and Xiao 2013). While most of the
57 methods in these previous studies rely exclusively on an ensemble-based error covariance,
58 several recent studies did combine the ensemble covariances with time-invariant static
59 covariances in their 4DEnVar, i.e. hybrid 4DEnVar (Buehner et al. 2013; Desroziers et al. 2014;
60 Lorenc et al. 2014; Wang and Lei 2014). Formulating the problem in the variational framework
61 allows one to take full advantage of the many developments that have taken place over the years,
62 such as dynamic constraints (Gauthier and Thépaut 2001; Kleist et al. 2009) and variational bias
63 correction (Derber and Wu 1998; Dee 2005; Zhu et al. 2014).

64 In Part I (Kleist and Ide 2014), it was demonstrated that including ensemble covariances
65 in a variational-based hybrid algorithm yielded improvements in the quality of analyses and
66 subsequent forecasts for the National Centers for Environment Prediction (NCEP) global

67 forecast system (GFS) model in the context of an observing system simulation experiment
68 (OSSE). The experiments were performed using 3DVar and hybrid 3DEnVar, leaving
69 significant room for improvement. Without access to the TL and AD models, a natural extension
70 of the hybrid 3DEnVar to include 4D ensemble perturbations (hybrid 4DEnVar), is a logical next
71 step for improving upon the previous work.

72 Several previous studies have investigated the use of 4DEnVar for use with global
73 deterministic NWP. A 4DEnVar algorithm was implemented into a prototype observation space
74 global data assimilation system for the United States Navy in Bishop and Hodyss (2011). There,
75 they focused on an adaptive localization algorithm within the context of a single case study.
76 Buehner et al (2010b) performed an inter-comparison study for the Canadian operational global
77 NWP model, and found that the 4DEnVar improved upon their operational, non-hybrid 4DVar in
78 the tropics and southern hemisphere, but not in the northern hemisphere. It was also found that
79 4DEnVar performed slightly worse than a hybrid 4DVar. Buehner et al. (2013), again using the
80 operational Canadian system, found that while the use of 4D instead of 3D ensemble covariances
81 did result in small, consistent improvements in their EnVar for deterministic NWP, the gains
82 were not as large as found when going from 3DVar to 4DVar. They also found that the
83 performance of their 4DEnVar was comparable to or better than their 4DVar, except for the
84 extratropical summer regions. Using the United Kingdom Met Office global system, Lorenc et
85 al. (2014) found that both hybrid 4DVar and 4DEnVar beat their 3D counterparts in NWP trial
86 runs. However, it was also found that their hybrid 4DEnVar did not perform as well as their
87 hybrid 4DVar, due to the heavy reliance on the climatological covariance that the 4DVar
88 algorithm can propagate through the assimilation window whereas the EnVar algorithm cannot.
89 Lastly, Wang and Lei (2014) performed a comparison study of hybrid 4DEnVar with 3DEnVar

90 using the NCEP GFS (model) at low resolution. They found that 4DEnVar was better than
91 3DEnVar, with a larger impact in the extratropical troposphere than in the tropics. They also
92 found that analysis increments from the 4DEnVar algorithm were more balanced than 3DEnVar,
93 and as in a companion 3DEnVar study (Wang et al. 2013), the use of a dynamic constraint was
94 valuable for improving the forecast skill of the 4DEnVar initialized forecasts in the extratropics.

95 This work is similar to the aforementioned studies that investigated the use of 4DEnVar
96 for initializing global deterministic NWP. Here, the NCEP GFS is utilized as in Wang and Lei
97 (2014), however, an OSSE is used instead of real observations as in Part I (Kleist and Ide 2014).
98 OSSEs have the distinction of allowing for the calculation of actual analysis error since the
99 “truth” is known. Furthermore, the experiments utilize a dual-resolution paradigm as in Part I.
100 Given the importance of initialization within the context of 4DEnVar for the NCEP GFS system
101 (Wang and Lei 2014) and the Met Office system (Lorenc et al. 2014), we also aim to corroborate
102 previous findings as well as explore alternate, computationally efficient initialization options.
103 The remainder of the manuscript is organized as follows. Section 2 describes the
104 implementation of the 4D extension to the hybrid including a time-invariant static error
105 covariance supplement. Section 3 then follows with a description of various OSSE-based
106 experiments that demonstrate the impact of utilizing 4DEnVar and hybrid variants relative to the
107 3D hybrid experiments that were carried out in Part I. Several experiments are carried out to
108 demonstrate the impact of including a static error covariance in the dual-resolution 4DEnVar
109 paradigm, as well as to show the impact of various dynamic constraints. A summary and
110 motivation for future work then follows.

111

112 2. 4D Extensions of 3D Hybrid

113 A. GSI-BASED HYBRID

114 Traditionally, incremental 4DVar involves solving for the optimal solution (\mathbf{x}'_0) at the
 115 beginning of a time window, obtained by minimizing a cost function

$$116 \quad J(\mathbf{x}'_0) = \frac{1}{2}(\mathbf{x}'_0)^T \mathbf{B}_f^{-1}(\mathbf{x}'_0) + \frac{1}{2} \sum_{k=1}^K (\mathbf{H}_k \mathbf{x}'_k - \mathbf{y}'_k)^T \mathbf{R}_k^{-1} (\mathbf{H}_k \mathbf{x}'_k - \mathbf{y}'_k) \quad (1)$$

$$118 \quad \mathbf{x}'_k = \mathbf{M}_k \mathbf{x}'_0 \quad (2)$$

120 where \mathbf{B}_f is the static background error covariance, \mathbf{R} the observation error covariance, \mathbf{H} the
 121 linearized observation operators, and \mathbf{y}' the observation innovations. Note that preconditioning
 122 is being ignored for simplicity. The solution is determined by using the asynchronous
 123 observations up to K -time levels and the TL model (\mathbf{M}_k) for each time-level index (k). The AD
 124 model (transpose of \mathbf{M}_k) is necessary in order to obtain the gradient for the minimization. This
 125 cost function can then be extended to a hybrid 4DVar cost function by including an ensemble
 126 control variable ($\boldsymbol{\alpha}^n$) for the ensemble contribution to the analysis increment at the beginning of
 127 the window:

$$128 \quad J(\mathbf{x}'_f, \boldsymbol{\alpha}) = \beta_f \frac{1}{2}(\mathbf{x}'_f)^T \mathbf{B}_f^{-1}(\mathbf{x}'_f) + \beta_e \frac{1}{2} \sum_{n=1}^N (\boldsymbol{\alpha}^n)^T \mathbf{L}^{-1}(\boldsymbol{\alpha}^n) +$$

$$\frac{1}{2} \sum_{k=1}^K (\mathbf{H}_k \mathbf{x}'_k - \mathbf{y}'_k)^T \mathbf{R}_k^{-1} (\mathbf{H}_k \mathbf{x}'_k - \mathbf{y}'_k) \quad (3)$$

129 where \mathbf{L} denotes the error covariance for the ensemble control variable, specified to be unit
 130 amplitude and Gaussian in structure, thereby acting to enforce localization of the ensemble based
 131 error covariance. Here, the localization is assumed to be static and does not evolve with the
 132 flow, which Fairbairn et al. (2014) demonstrates to be potentially detrimental to 4DVar

133 relative to 4DVar using a toy model. Just as in the 3D hybrid case, \mathbf{x}'_0 becomes a linear
 134 combination of that which is derived from a static error covariance (\mathbf{x}'_f) and that which is
 135 derived from the ensemble perturbations (\mathbf{x}_e^n) as prescribed by the control variable ($\boldsymbol{\alpha}^n$) for an
 136 ensemble of size N :

$$137 \quad \mathbf{x}'_0 = \mathbf{x}'_f + \mathbf{T} \sum_{n=1}^N (\boldsymbol{\alpha}^n \circ (\mathbf{x}_e)_0^n) \quad (4)$$

138 Here, both \mathbf{x}'_f and $(\mathbf{x}_e)_0^n$ are valid at $t=0$ and the hybrid analysis increment is propagated to each
 139 time level (k) by the TL as in Eq. (2). For a dual-resolution configuration, \mathbf{T} represents the
 140 interpolation (or other manipulation) from ensemble to deterministic resolution. This hybrid
 141 configuration will hereafter be referred to as H-4DVar, recalling that it requires the AD (and
 142 forward TL) model within the minimization.

143 As pointed out in previous studies (Buehner 2013; Desroziers et al. 2014; Lorenc et al.
 144 2014; Wang and Lei 2014), this can be further manipulated to solve a similar 4D cost function as
 145 in (4) without the need for the TL and the AD by utilizing 4D ensemble perturbations (hybrid
 146 4DEnVar), where the increment valid at each of the observation time levels can be prescribed as

$$147 \quad \mathbf{x}'_k = \mathbf{M}_k \mathbf{x}'_f + \mathbf{T} \sum_{n=1}^N (\boldsymbol{\alpha}^n \circ (\mathbf{x}_e)_k^n). \quad (5)$$

148 For hybrid 4DEnVar, \mathbf{M}_k is simply chosen to be the identity model, where a single static
 149 contribution is valid through the whole window as in hybrid 3DEnVar. The ensemble
 150 contribution to the increment now utilizes 4D, nonlinear perturbations associated with each of
 151 the observation bins. In this particular formulation, $\boldsymbol{\alpha}^n$ is assumed to be the same throughout the
 152 assimilation window, analogous to the weights in a 4D-LETKF without temporal localization.
 153 More details regarding the GSI 4DEnVar formulation can be found in Wang and Lei (2014).
 154

155 Alternate formulations and practical implementations of 4DEnVar are provided in Desroziers et
156 al. (2014).

157 To allow for a single time level, the time invariant contribution, \mathbf{x}'_f is simply chosen to
158 be prescribed in the center of the window and \mathbf{M}_k is set to be an identity model. One could
159 formulate an alternate hybrid and utilize the TL and AD models to evolve the static contribution
160 *only*, prescribe \mathbf{x}'_f to be at the beginning of the window, and linearly combine with the 4D
161 ensemble contribution, as in Eq. (5). For the purposes of this work, only the former possibility
162 where the identity model is used and the static contribution is defined to be valid at the center of
163 the window is considered due to the computational advantage. The special case where the static
164 contribution is set to zero ($\beta_f^{-1}=0$) will be referred to as 4DEnVar, whereas any configuration
165 that utilizes a time-invariant static contribution will be referred to as hybrid 4DEnVar (or H-
166 4DEnVar). Given a configuration that utilizes a dual-resolution option, the static contribution is
167 aimed at accomplishing two things: 1) filling the null space at higher spatial frequencies that the
168 low resolution ensemble cannot resolve, assuming there are observations at such frequencies, and
169 2) providing more search directions for the minimization, complementing the relative few such
170 search directions available via localized linear combinations of the ensemble perturbations.

171 The necessary components to solve the GSI using the formulation that is described by
172 Eqs. (3) and (5), 4DEnVar, are already in place. The ensemble control variable for the hybrid is
173 developed and implemented as part of the 3D-hybrid work described in Part I. The 4D ensemble
174 perturbations are derived from a serial square root filter form of an ensemble Kalman filter
175 (EnKF; Whitaker and Hamill 2002; Whitaker et al. 2008). A 4DVar capability within the GSI
176 was previously developed through collaboration with colleagues at NASA GMAO (Todling and
177 Trémolet, personal communication). A key component of the extension of 3DVar to 4DVar

178 within the GSI was the addition of new observation handling features and the inclusion of time
179 binning. All configurations (3D and 4D) utilize similar default observation selection. With these
180 pieces in place, the 3D hybrid capability is extended to allow for 4D ensemble perturbations and
181 a 4DEnVar option.

182

183 B. SINGLE OBSERVATION EXAMPLE

184 To demonstrate the impact of the 4DEnVar algorithms, a set of experiments that
185 assimilate only a single simulated observation is performed. A summary and description of the
186 experiments is available in Table 1. The two hybrid cases that include a static and ensemble
187 contribution to the increment are done so with $\beta_f^{-1}=0.25$ and $\beta_e^{-1}=0.75$. For the two 4DVar
188 variants, the TL and AD that are utilized in the inner loop are adaptations of a simple model that
189 has been under development at NCEP for the sole purpose of running 4DVar. Although not yet
190 ready for large problems or cycling due to its current computational inefficiencies, it is still quite
191 useful for evaluating low-resolution single observation experiments that require only a few
192 iterations to minimize. For computational reasons, a single resolution of T126 with a 40 member
193 ensemble is utilized for the single observation example for demonstration. All of the solutions
194 that have an ensemble contribution to the increment use Gaussian localization with a cutoff of
195 1200 km in the horizontal and 1.2 scale heights in the vertical. The single observation chosen for
196 the experiments is a mid-tropospheric temperature observation taken at the beginning of a 6-hr
197 assimilation window, three hours prior to the “update” analysis time, and is assigned a negative
198 two degree departure from the background at that time (a 3-hr forecast) and a one degree
199 observation error standard deviation. The observation is chosen to reside upstream from the base

200 of a shortwave trough, similar to the single observation experiments shown in Buehner et al.
201 (2010a) for comparison.

202 The resultant analysis increment at the middle of the assimilation window, three hours
203 after the observation is taken, for the various 4D configurations is shown in Fig. 1. All four
204 experiments show the maximum increment downstream from where the observation was taken
205 consistent with the northwesterly background flow. The 4DVar experiment results in a spatially
206 broad, quasi-Gaussian temperature increment (Fig. 1, upper left). This is not terribly surprising
207 given that only three hours elapse between the time that the observation was taken at the
208 beginning of the window and the analysis time at the center of the window. The three
209 experiments that utilized ensemble covariances exhibit a temperature increment that is stretched
210 along the height gradient as would be expected (Fig. 1). All of the experiments show a cyclonic
211 wind response to the cold temperature observation and increment, with the ensemble- and
212 hybrid-based experiments showing a stronger wind response than the 4DVar case. It is clear that
213 the 4DEnVar case suffers from sampling (spurious correlations) more so than the two hybrid
214 variants. This is not surprising either, in this case, given that a 40-member ensemble is utilized
215 in combination with quite broad localization. However, adding a time-invariant, static
216 contribution to the increment helps to reduce the impact of these apparent problems substantially,
217 without hurting the 4D nature of the increment. In fact, the H-4DEnVar increment is
218 qualitatively and quantitatively very similar to that from the H-4DVar experiment, despite the
219 fact that one uses a fairly simply dynamic model while the other utilizes 4D, nonlinearly evolved
220 ensemble perturbations. This is in contrast to the single observation experiments carried out in
221 Lorenc et al. (2014), where the more heavily reliance on a static, time-invariant contribution to
222 the increment was shown to be problematic in the H-4DEnVar paradigm.

223 The same single observation test can be utilized to further investigate how the algorithms
224 handle the propagation of information through the assimilation window, by visualizing the
225 analysis increment at various k -time levels. Recall that for the 4DVar cases, this involves an
226 explicit propagation of the increment through the use of a linear, dynamic model, i.e. Eq. (2).
227 For the 4DEnVar variants, the propagation of information is achieved implicitly through
228 correlations contained within the 4D ensemble perturbations, i.e. Eq. (5) with $\mathbf{M}_k=\mathbf{I}$. The time
229 evolution of the temperature and wind increments at 500 hPa in 3-hr intervals for the H-4DVar
230 and H-4DEnVar cases is shown in Fig. 2. First, note that the solutions of the two experiments
231 are identical at the beginning of the window, the time at which the observation was taken. This
232 is expected given the same ensemble perturbations, relative weights, and static error covariance,
233 since the TL and/or nonlinear models have not evolved the increments in the two experiments.
234 This is not the case for observations taken at other times in the assimilation window (not shown),
235 where for the H-4DVar case, the effective “static” error covariance is evolved forward in time
236 consistent with the linear dynamic model. For other times in the six hour assimilation window,
237 middle and end respectively, the temperature and wind increments for the two hybrid cases are
238 strikingly similar both qualitatively and quantitatively. By the end of the window, the H-
239 4DEnVar case has more small amplitude features away from the location of the original
240 observation and has a slightly tighter structure aligned with the height contours. The H-4DVar
241 case, on the other hand, has somewhat broader features at the end of the window. It is quite
242 encouraging that the computationally cheaper ensemble method can mimic quite nicely the
243 evolution of the increment within the assimilation window. Although not shown, other single
244 observation experiments for other variables and levels reveal similar results. A similar
245 comparison for different single observation tests can be found in Lorenc et al. (2014).

246 3. Constraints on high frequency noise

247 The 4D_{En}Var (and hybrid) option is implemented in such a way as to allow for the
248 application of many of the standard features included in the 3D GSI such as variational quality
249 control, variational satellite bias correction, the tangent linear normal mode constraint (TLNMC;
250 Kleist et al. 2009), and various weak constraints such as digital filter (Gustafsson 1993;
251 Polavarapu et al. 2000; Gauthier and Thépaut 2001). However, the 4D_{En}Var solution to the
252 analysis problem requires special attention when it comes to the application of such constraints
253 given that the 4D aspect of the problem is obtained through a dot product of the weights and the
254 ensemble perturbations in a discrete manner without the explicit use of a dynamic model. The
255 TLNMC and the weak constraint digital filter, as well as the combination of the two, are
256 explored within the context of 4D_{En}Var. In this work, the constraints are imposed on the high-
257 resolution deterministic analysis. While the EnKF analysis is re-centered about the filtered,
258 deterministic analysis, the individual members are only explicitly filtered using the full-field
259 digital filter of the GFS model.

260

261 A. TANGENT LINEAR NORMAL MODE CONSTRAINT

262 For a dynamic constraint such as the TLNMC, the minimization procedure of the cost
263 function within the 4D context can be prohibitively expensive if one were to use hourly time
264 levels and a six-hour assimilation window, as an example. Furthermore, the original
265 implementation of the normal mode constraint for application within hybrid 3D_{En}Var requires
266 the filter to be applied to the total analysis increment as a sum of static and ensemble
267 contributions without much flexibility, as in Eq. (2) from Part I (Kleist and Ide 2014). For these

268 reasons, several options related to the application of various constraints are explored and
 269 implemented into the analysis code.

270 Building on previous successes in implementing the TLNMC in the 3DVar (Kleist et al.
 271 2009) and hybrid 3DEnVar (Part I) context, various possibilities for application of such a
 272 constraint in the 4DEnVar and hybrid context are developed and tested. The first possibility is to
 273 apply the TLNMC to only the static contribution to the increment:

$$274 \quad \mathbf{x}'_k = \mathbf{C} \mathbf{x}'_f + \mathbf{T} \sum_{n=1}^N \left(\boldsymbol{\alpha}^{n \circ} (\mathbf{x}_e)_k^n \right). \quad (6)$$

275 This requires a single constraint operation per iteration, making the cost equivalent to 3D
 276 applications. Note that this capability can also be applied within the 3D hybrid context, though
 277 previous results demonstrate that the inclusion of the constraint to the total increment aids in
 278 reducing potential imbalances introduced by the ensemble based increment (Wang et al. 2013;
 279 Wang and Lei 2014). This formulation, given by Eq. (6), is motivated by the desire to allow for
 280 the use of the unfiltered ensemble covariance while still filtering the static contribution to the
 281 increment.

282 In order to help mitigate imbalances that can be introduced through sampling and
 283 localization, two additional possibilities to apply the TLNMC to the total increment are explored
 284 in the H-4DEnVar (and 4DEnVar) context. One is to simply apply the constraint over all k -time
 285 levels:

$$286 \quad \mathbf{x}'_k = \mathbf{C}_k \left[\mathbf{x}'_f + \mathbf{T} \sum_{n=1}^N \left(\boldsymbol{\alpha}^{n \circ} (\mathbf{x}_e)_k^n \right) \right]. \quad (7)$$

287 Here, the balance operator \mathbf{C}_k is also denoted with a time level index, since the possibility exists
 288 for linearization about the background for each time level, as the constraint is in fact tangent-
 289 linear. However, for practical reasons such as computational cost and memory, the background
 290 state used for linearization in \mathbf{C} is taken from the center of the observation window and applied

291 to all k time levels. Wang and Lei (2014) utilized the constraint in this manner for their
 292 4DEnVar experiments and found improvements in forecast skill in the extratropical troposphere.
 293 Even still, the application of the constraint in this manner can be prohibitively expensive. For a
 294 six-hour window with hourly time levels and hourly ensemble perturbations, the computational
 295 cost of the analysis goes up substantially. For this reason, another possibility to apply the full
 296 constraint only to the solution in the middle of the window is considered. The advantage of such
 297 a method is that the increment that is applied to the background and used to restart the model can
 298 be filtered explicitly reducing spin-up and spin-down, without the considerable cost of having to
 299 filter all time levels. This of course introduces inconsistencies between the incremental solution
 300 in the center of the window and the other time-levels, which could have undesirable
 301 consequences. All applications of the TLNMC referred to hereafter utilize the default 8 vertical
 302 modes and a single iteration as in Kleist et al. (2009) and Part I (Kleist and Ide 2014).

303

304 B. WEAK CONSTRAINT DIGITAL FILTER

305 Digital filter weak constraints have been shown to be beneficial in many 4DVar
 306 applications (Polavarapu et al. 2000; Gauthier and Thépaut 2001; Wee and Kuo 2004). A similar
 307 constraint has previously been implemented by colleagues at NASA GMAO for the 4DVar
 308 extension of the GSI.¹ We denote this digital filter initialization as the JCDFI. The formulation
 309 of the JCDFI involves the addition of a new penalty term (see Gustafsson 1993 or Polavarapu et
 310 al. 2000 for a more detailed derivation):

$$311 J_{\text{dfi}} = \chi \langle \mathbf{x}_m - \mathbf{x}_m^i, \mathbf{x}_m - \mathbf{x}_m^i \rangle \quad (8)$$

¹ The JCDFI that was originally developed for 4DVar applications was not implemented for the standard double conjugate gradient minimization algorithm that is utilized for the hybrid applications. Minor modifications were necessary to adapt the code for use in this context.

312 where \mathbf{x}_m^i denotes a filtered or “initialized” state at time m (in the case of 4DEnVar, assumed to
 313 be the center of the assimilation window), and χ a general weighting factor. This weighting
 314 parameter is typically denoted α in the literature but avoided in this context to remove confusion
 315 with the ensemble control variable. The filtered state is constructed from the 4D increment using
 316 the same filter coefficients (h) as in the standard formulation (Lynch and Huang 1992)

$$317 \quad \mathbf{x}_m^i = \sum_{k=1}^K h_{k-m} \mathbf{x}_k^u \quad (9)$$

318 where \mathbf{x}_k^u denotes the unfiltered increment at each time level- k . The norm for the penalty
 319 function in Eq. (8) is chosen to be a dry energy norm, though the capability does exist to use a
 320 moist energy norm. Such a constraint in the 4DEnVar context, on one hand, potentially allows
 321 for noise control within the 4D increment without adding much computational cost. On the other
 322 hand, high frequency noise cannot be described by the hourly incremental states in this context.
 323 Furthermore, there are no means to correct the origin of the high frequency noise with links to
 324 the actual model equations as would be the case for a 4DVar-based formulation. Despite these
 325 limitations, it does allow for a cost effective alternative to employ some filtering of the 4D
 326 increment. Additionally, it is possible to explore the use of a combination of noise constraints,
 327 weak constraint digital filter and TLNMC, to improve the quality of analysis.

328

329 C. SINGLE ANALYSIS IMPACT

330 To ensure that the constraints are all properly functioning for use with the 4DEnVar
 331 paradigm within GSI, a single analysis test case that assimilates real observations including
 332 satellite radiances within a 6-hr window valid at 06 UTC 15 July 2010 is performed. The
 333 background and ensemble members are generated from an offline experiment that utilizes the
 334 dual-resolution hybrid configuration, but at higher resolution than the OSSE-based cycled

335 experiments described in Part I, with T574 for the high-resolution deterministic component and
336 T254 for the 80-member ensemble. Double inner-loops of 100 iterations each are utilized,
337 whereby the quality control, observation selection, and re-linearization are performed and the
338 nonlinear model is not rerun. The analysis is run with four separate configurations: 1) 4DEnVar,
339 2) 4DEnVar with TLNMC on the increment over all time levels, 3) 4DEnVar with JCDFI, and 4)
340 4DEnVar with the TLNMC imposed at the center of the window only combined with the JCDFI.
341 The fourth configuration that utilizes both constraints is done with computational considerations
342 in mind. In all four configurations, there is no static contribution to the solution, i.e., $\beta_f^{-1}=0$.
343 The desire is to have the TLNMC remove gravity mode tendencies in the middle of the window
344 while allowing the weak constraint digital filter to maintain the connection to the other time
345 levels through the procedure to get the filtered state, and penalize additional noise accordingly,
346 without having to use the normal mode procedure explicitly over all time levels.

347 The divergence increment for the experiments that ran with the JCDFI is damped across
348 the entire spectrum, including at large scales (Fig. 3). Some interesting behavior is observed on
349 the high spatial frequency part of the spectrum, most of which is an artifact of the dual-resolution
350 aspect of the configuration. The aliasing that results from the interpolation within the current
351 dual-resolution configuration is an area of active research and the subject of a future manuscript.
352 As expected, the more the solution is constrained, the more difficult it becomes to draw to the
353 observations (Fig. 4). At the end of the first inner loop minimization (100 iterations), the penalty
354 reduction is greatest for the 4DEnVar case with no extra constraints, followed by the
355 experiments that utilized a single constraint (TLNMC and JCDFI), followed by the configuration
356 that used both constraints. At the end of the second inner loop minimization, the order is
357 slightly changed, but the final observation penalty is within one percent for all four

358 configurations. The minimization appears well behaved with a steady reduction of the gradient
359 norm in all four experiments. By design, the TLNMC has a significant impact on the
360 incremental tendencies (Kleist et al. 2009), particularly the gravity mode tendencies (Table 2).
361 On the other hand, while the JCDFI reduces the total tendencies by a factor of two, it has very
362 little impact on the amount of tendencies that project onto undesirable gravity modes due to the
363 aforementioned issues regarding the inability to filter the highest frequency noise using only
364 hourly states. The combined (COMB) configuration that utilizes JCDFI and TLNMC in the
365 center of the window appears to be the best compromise, by significantly reducing the total
366 tendencies while also reducing the percentage that project onto gravity modes, all at a
367 significantly cheaper computational cost than running the normal mode constraint on all time
368 levels. The inclusion of a static error covariance (H-4DEnVAR) also acts to decrease the
369 incremental tendencies slightly while improving the ratio of gravity mode to total tendencies.

370 4. Cycled Experiment Results

371 A. ANALYSIS ERROR COMPARISON

372 A series of experiments is designed to study the impact of the 4D-ensemble and hybrid
373 components as an extension to the 3D experiments that were described in Part I. All of these
374 new experiments utilize the same cold start initial condition for the control and ensemble,
375 simulated observations from the Joint OSSE (Andersson and Masutani 2010), GFS model
376 version and spatial resolutions (dual resolution mode, T382 deterministic with an 80 member
377 T190 ensemble), comparable versions for the GSI and EnKF codes, identical inflation
378 parameters (Whitaker and Hamill 2012), and cycling configuration for the data assimilation
379 including the re-centering procedure. The experiments are designed in such a way as to infer
380 which components of the hybrid and 4D extensions yield improvements (or degradations) to the

381 quality of analyses. A listing and description of the experiments can be found in Table 3. Note
382 that the two 3D experiments were carried out for Part I, but are included here as a reference
383 point. To reiterate, the assimilation and nature run models are not the same, and it is important
384 to keep in mind their different handling of various components such as parameterized convection
385 when interpreting the results.

386 The first experiment, H-4DENVAR_NMI, is designed to build upon the successes of the
387 3DHYB experiment in Part I and to test the impact of going from 3D to 4D ensemble
388 perturbations. The assimilation window is kept fixed at six hours, though the H-
389 4DENVAR_NMI experiment utilizes hourly time levels and therefore, hourly background and
390 ensemble perturbations, e.g., Eqs. (3), (5), and (7). The observation selection procedure is
391 identical in both configurations. This hybrid configuration also utilizes a 25% time-invariant
392 static contribution to the analysis increment as well as the TLNMC over all seven time levels.
393 The change in the analysis error for zonal wind, temperature, and specific humidity is shown in
394 Fig. 5. Generally speaking, the analysis error is smaller in the H-4DENVAR_NMI experiment,
395 especially for upper tropospheric extratropical winds and temperature, and lower tropospheric
396 water vapor. It appears that by going from 3D to 4D, the temperature analysis error has actually
397 increased over the southern polar cap in the lower troposphere. Also of note is the increase in
398 analysis error for specific humidity in the mid-tropospheric tropics, temperature in the upper
399 tropospheric tropics, as well temperature in the upper stratosphere. The increased analysis error
400 is probably related to differences in topography or model physics, as well as the use of three
401 hourly output of the nature run to generate the continuous simulated observation set (Errico et al.
402 2012). In particular, differences between the nature run model and assimilation model in terms
403 of the handling of vertical mixing in very stable regimes, deep convection in the tropics, and

404 damping in the upper stratosphere are likely being exposed. Other 4D experiments exhibit
405 similar behavior. By in large, however, the impact in going from 3D to 4D is generally positive,
406 consistent with the better use of observations distributed through the assimilation windows as
407 was demonstrated with the single observation test. However, the impact is smaller than what
408 was found in going from 3DVar to hybrid 3DEnVar (Part I, Kleist and Ide 2014), consistent with
409 the findings of Buehner et al. (2013) and not as impressive as the findings in Wang and Lei
410 (2014).

411 Although encouraged by the fact that the H-4DEnVar_NMI yielded improved analyses
412 relative to the original 3D hybrid experiment, the application of the TLNMC over all seven
413 observation bins is computationally expensive, given the necessity to calculate incremental
414 tendencies as well as grid to spectral transforms, all within the iterative scheme. An experiment
415 is carried out to test the impact of replacing the TLNMC with a more computationally efficient
416 JCDFI (H-4DEnVar_DFI). The digital filter term is based on a dry energy norm and utilizes a
417 weighting parameter (χ) of 10 in Eq. (8), based on previous findings (Polavarapu et al. 2000;
418 Gauthier and Thépaut 2001). Relative to an experiment that does not utilize any of the constraint
419 options (H-4DEnVar), the H-4DEnVar_DFI is found to have small, consistently negative impact
420 (Fig. 6, purple). Sensitivity experiments that varied the amplitude of the weak constraint cost
421 function parameter did not yield significant improvements, as a large increase in the parameter
422 amplitude results in a significantly worse fit to observations.

423 Similar to the findings of Kleist et al. (2009), the TLNMC does have an impact reducing
424 the background and analysis errors for surface pressure whereas the JCDFI does not (Fig. 7). In
425 fact, the H-4DEnVar_DFI results in increased background error for surface pressure relative to
426 no constraint for many cycles (Fig. 7 top). The JCDFI as designed is not able to filter out higher

427 frequency noise, given that it cannot be described with a time series of hourly states (Sec. 3c).
428 Furthermore, there is no direct link to the model equations themselves as there is in the 4DVar
429 variant of the digital filter. In contrast, the H-4DEnVar_NMI results in a consistent reduction of
430 analysis error relative to the experiment with no constraint (Fig. 6, blue), consistent with the
431 findings of Wang et al. (2013), Kleist and Ide (2014), and Wang and Lei (2014).

432 Despite these shortcomings, an experiment is carried out that utilizes a combined
433 constraint similar to that which was utilized in section 3. In this new experiment (H-
434 4DEnVar_COMB), the same digital filter parameters from H-4DEnVar_DFI are utilized while
435 the TLNMC is only applied to the center of the assimilation window, or in other words, to the
436 increment that is added to the 6-hr forecast which is then passed on as the initial condition for the
437 forecast model. The JCDFI uses information from all time levels, and therefore aids to remove
438 some gravity mode tendencies for the time levels away from the center according to the filter
439 weights without having to explicitly apply the normal mode initialization over all time levels.
440 This is demonstrated by looking at the total and gravity mode tendencies from the single analysis
441 case used in section 3. The largest incremental tendencies are generated for the analysis that
442 utilized no constraint whereas the TLNMC is most efficient at reducing them (Fig. 8 top).
443 While the JCDFI does act to reduce the incremental tendencies, it also acts to increase the
444 amount of gravity mode tendencies that project onto the total tendencies (Fig. 8 bottom). What
445 is especially troubling is the fact that the JCDFI increases the ratio of gravity mode to total
446 tendencies to be larger than with no constraint at all. The combined filter is most efficient at
447 reducing the incremental tendencies at the center of the window as expected, but the ratio of
448 gravity mode to total tendencies is not as small as the TLNMC-only over all time levels. Like
449 the JCDFI itself, the combined constraint acts to degrade the quality of analyses for most levels,

450 the exception being the layer between 300 hPa and 20 hPa (Fig. 6). Additionally, the combined
451 constraint is worse than the multi time-level TLNMC for all layers except for a slight
452 improvement found between 300 hPa and 150 hPa. It is clear from this set of experiments that of
453 the constraint formulations considered, the TLNMC is the only one that consistently yields
454 improved analyses.

455 The suite of constraint experiments all included a 25% contribution from the static error
456 covariance. It was found in Part I that including a static **B** contribution had very little impact on
457 the solution. To explore this further, an additional 4D_{En}Var experiment is run to test the impact
458 of a static contribution in the absence of other constraints. Surprisingly, even though the static
459 contribution is time invariant, it does act to reduce analysis error in the 4D hybrid context (Fig.
460 9). The discrepancy between this result and the findings in Part I and Wang et al. (2013) and
461 Wang and Lei (2014) can be attributed to the use of the TLNMC and use of a dual-resolution
462 configuration. The TLNMC is in fact improving upon the quality of the hybrid analyses even
463 further (Fig. 6). However, since the TLNMC is actually a strong constraint on the incremental
464 solution, the inclusion of a static contribution may be unnecessary for certain configurations.
465 This encouraging result provides motivation to explore the use something like the TLNMC
466 within the EnKF analysis update.

467

468 B. FORECAST IMPACT

469 Similar to Part I, a forecast impact experiment is carried out utilizing 00 UTC analyses
470 from the H-4D_{En}Var_NMI configuration to initialize the GFS model for integration out to 7.5
471 days. The first two weeks of the experiment are ignored to account for spin-up. All verification
472 is done relative to the ECMWF nature run on a common grid. It was shown in Part I that

473 forecasts initialized from the 3DHYB analyses were generally superior to those that were
474 initialized from 3DVar analyses. The analyses from the H-4DEnVar_NMI experiment are
475 chosen in order to directly compare to the 3D hybrid experiments carried out in Part I, where the
476 TLNMC was also utilized, in addition to the fact that the H-4DEnVar_NMI experiment resulted
477 in the smallest analysis errors relative to the other 4D hybrid experiments.

478 A comparison of the 500 hPa geopotential height anomaly correlation die off curves
479 reveals that the 4D hybrid experiment yields improved forecasts at this level for all lead times in
480 both extratropical hemispheres (Fig. 10). The improvement over the 3D hybrid experiment is not
481 as large in amplitude as was found in going from 3DVar to hybrid 3DEnVar. The 4D hybrid
482 configuration results in improved height forecasts at other levels beyond 72-h, except for the
483 upper stratosphere (Fig. 11). This is consistent with the increased analysis temperature errors
484 that were found for these upper levels (Fig. 5 middle), related to the ensemble under-dispersion
485 in terms of temperature in the upper stratosphere (noted in Part I).

486 The forecast impact on other variables and levels is more mixed (Fig. 12), similar to the
487 analysis error findings (Fig. 5). In general, it can be seen that the only region that resulted in a
488 consistent improvement in going from 3D to 4D hybrid was in the northern hemisphere
489 troposphere (Fig. 12, left column). While there are some improvements in the wind and height
490 forecasts in the southern hemisphere beyond 48 hours, the differences are generally not
491 statistically significant. The impact in the tropics is generally found to be neutral to worse. This
492 is counter to the findings in Part I in going from 3DVar to 3D hybrid (Kleist and Ide 2014),
493 where significant improvements were found in the tropics and southern hemisphere troposphere.
494 The use of 4D versus 3D ensembles results in only modest improvements, consistent with the
495 findings in Buehner et al. (2013). This is somewhat counter to the findings of Wang and Lei

496 (2014), where large improvements were found in their 4D_{EnVar} versus 3D_{EnVar} in the
497 extratropics, particularly in the Southern Hemisphere. However, they too found only very small
498 impacts in the tropics.

499 5. Summary and Conclusions

500 An extension of the GSI-based hybrid variational ensemble algorithm to include 4D
501 ensemble perturbations is proposed and implemented. The 4D_{EnVar} algorithm has several
502 advantages relative to 4D-EnKF and 4DVar algorithms that make it attractive for an operational
503 center such as NCEP. Since 4D_{EnVar} does not require the use of additional dynamic models
504 (i.e., the TL and AD), it requires less in terms of computational resources than a traditional
505 4DVar algorithm. Since it is based on a variational algorithm, it becomes quite trivial to
506 supplement the ensemble-based (4D) covariance with some static estimate that can help
507 ameliorate potential sampling issues, particularly for small ensemble sizes. Furthermore, such a
508 configuration allows for easy implementation of a dual-resolution algorithm, solving for a high-
509 resolution increment using a low-resolution ensemble and control variable (or vice versa,
510 analogous to the interpolation of ensemble weights algorithm, Yang et al. 2009). Lastly,
511 variational-based methods have the advantage of physical space localization, implicit through the
512 localization of the ensemble control variable, which can be important for observations such as
513 satellite radiances.

514 Several OSSE-based experiments are carried out to demonstrate that the 4D_{EnVar}
515 algorithm contributes positively to the quality of analysis relative to the 3D_{EnVar} and 3DHYB
516 experiments that were carried out in Part I. It is found that going from 3D to 4D ensemble
517 perturbations reduces the analysis error for most variables and levels, though the impact is
518 generally small as in Buehner et al. (2013), in contrast to the findings for the extratropics in

519 Wang and Lei (2014). The addition of a time-invariant static contribution to the analysis
520 increment (i.e., hybridization) does improve the quality of analyses, but only in the absence of
521 the TLNMC. Experiments that utilized the JCDFI, either alone or in combination with the
522 TLNMC, resulted in increased analysis error due to the use of hourly states in prescribing the
523 filtered state. Despite the computational cost associated with it, the use of the TLNMC over all
524 4D time levels produced the highest quality analyses. Wang and Lei (2014) also found that the
525 use of the TLNMC improved forecast quality within their 4D_{EnVar} experiments. It is possible
526 that the JCDFI may still have utility if one were to utilize higher temporal frequency for the
527 observation bins and ensemble time levels. This will prove to be computationally expensive, but
528 may be worth pursuing as computing power continues to increase.

529 The H-4D_{EnVar}_NMI experiment had the largest impact in the extratropical troposphere.
530 This corresponds to a more dynamically active region, where the propagation of information and
531 the usage of the appropriate time levels within the observations are more important. The
532 amplitude of the error reduction in going from the 3DHYB to H-4D_{EnVar}_NMI is found to be
533 smaller than that when going from the 3DVar control to 3DHYB. Going from 3D to 4D in this
534 context resulted in increased analysis error for a few variables in certain regions. This could be
535 the result of the use of discrete (3 hourly) time levels from the nature run to generate the 4D
536 observations as well as differences between the assimilation and nature run models. Perhaps the
537 use of stochastic physics in place of additive inflation (and a single deterministic model) within
538 the ensemble could reduce the negative impact in going to 4D. There is work underway to
539 generate new high-resolution nature runs with better temporal output. It will be interesting if
540 some of these issues can be reduced with the new data sets. Lastly, it is found that the improved

541 analyses generally result in improved forecast skill, especially for extratropical tropospheric
542 height and wind forecasts.

543 Although the GSI has a 4DVar capability within it, the inefficiencies of the inner loop
544 dynamic model make it unaffordable for running fully cycled experiments. Work is ongoing to
545 make the dynamic model more efficient. Once it is ready and can be utilized for such an
546 experiment, it will be interesting to compare the results of the H-4DVar experiments with
547 actual 4DVar, including H-4DVar.

548 There is also additional work that can be done on the 4DVar algorithm itself, such as
549 treating the outer loop more appropriately, along the lines of the ideas proposed in Yang et al.
550 (2012). To this point, the 4DVar algorithm has been executed more like a 3DVar
551 implementation at least in terms of how the multiple inner loops operate. Perhaps some of the
552 noted issues can be improved through a proper outer loop where the nonlinear control model is
553 rerun but the perturbations are held fixed. An alternative would be to allow for the full
554 reintegration of the control and the ensemble, much like running in place (Yang et al. 2012) or an
555 iterative ensemble Kalman smoother (Bocquet and Sakov 2014). Lastly, work is underway to
556 explore the use of a 4D incremental analysis update (IAU, Bloom et al. 1996) within the hybrid
557 4DVar context. The 4DIAU has been demonstrated to be a suitable alternative to a weak
558 constraint digital filter for 4DVar (Lorenc et al. 2014). The 4DIAU is especially attractive as
559 it provides a means for passing the 4D incremental state to the forecast model, instead of the
560 single incremental state in the center of the assimilation window as is typically done.

561 In addition the aforementioned work related specifically to 4DVar, there is related
562 work that could be performed to improve the system. The TL and AD models that are used as
563 part of the TLNMC and 4DVar applications are still based on a dry, adiabatic tendency model.

564 This has been shown to have a negative impact in the tropics (Kleist et al. 2009; Wang et al.
565 2013; Wang and Lei 2014). Work is ongoing to add various moist physics to these models to
566 help improve the use of the TLNMC in the tropics. It should also be pointed out that the
567 selection of weighting parameters between the ensemble and static contribution has been
568 simplistic to this point, and could certainly be investigated further as in Bishop et al. (2013a,b).
569 One proposed modification to the application of the weighting already underway would be to
570 consider scale-dependent parameters (Chapter 4 of Kleist 2012).

571

572

573

574 *Acknowledgements:* This work was completed as part of the first author's Ph.D. thesis at the
575 University of Maryland-College Park, made possible through the support of EMC management,
576 especially Steve Lord, John Derber, and William Lapenta. The ECMWF nature run was
577 provided by Erik Andersson through arrangements made by Michiko Masutani. We are grateful
578 to the Joint Center for Satellite Data Assimilation for providing access to their supercomputers to
579 perform the experiments. The support of ONR grants (N000140910418, N0001410557) is
580 gratefully acknowledged. The authors wish to thank Nikke Privé and Ron Errico for kindly
581 providing access to the calibrated, simulated observations. The authors wish to thank Dave
582 Parrish and Jeff Whitaker for contributing to various aspects of the hybrid, EnVar, and EnKF
583 developments. Lastly, the authors acknowledge Ting Lei and Xuguang Wang for good
584 discussion and collaboration as they worked on an implementation of 4DEnVar in parallel to this
585 work. Two anonymous reviewers helped to significantly improve the manuscript.

586

587 **References**

- 588 Andersson, E., and M. Masutani, 2010: Collaboration on observing system simulation
589 experiments (Joint OSSE), ECMWF Newsletter No. 123, Spring 2010, 14-16. [Available
590 online at <http://www.ecmwf.int/publications/newsletters/pdf/123.pdf>]
- 591 Bishop, C. H. and D. Hodyss, 2011: Adaptive ensemble covariance localization in ensemble 4D-
592 Var state estimation. *Mon. Wea. Rev.*, **139**, 1241–1255.
- 593 Bishop, C. H., Elizabeth A. Satterfield, 2013a: Hidden Error Variance Theory. Part I: Exposition
594 and Analytic Model. *Mon. Wea. Rev.*, **141**, 1454–1468.
- 595 Bishop, C. H., Elizabeth A. Satterfield, Kevin T. Shanley, 2013b: Hidden Error Variance Theory.
596 Part II: An Instrument That Reveals Hidden Error Variance Distributions from Ensemble
597 Forecasts and Observations. *Mon. Wea. Rev.*, **141**, 1469–1483.
- 598 Bloom, S. C., L. L. Takacs, A. M. da Silva, D. Ledvina, 1996: Data Assimilation Using
599 Incremental Analysis Updates. *Mon. Wea. Rev.*, **124**, 1256–1271.
- 600 Bocquet, M. and P. Sakov, 2014: An iterative ensemble Kalman smoother. *Quart. J. Roy.*
601 *Meteor. Soc.*, **140**, 1521-1534.
- 602 Buehner, M., P. L. Houtekamer, C. Charette, H. L. Mitchell, and B. He, 2010a: Intercomparison
603 of variational data assimilation and the ensemble Kalman filter for global deterministic
604 NWP. Part I: Description and single-observation experiments. *Mon. Wea. Rev.*, **138**, 1550-
605 1566.
- 606 Buehner, M., P. L. Houtekamer, C. Charette, H. L. Mitchell, and B. He, 2010b: Intercomparison
607 of variational data assimilation and the ensemble Kalman filter for global deterministic
608 NWP. Part II: One-month experiments with real observations. *Mon. Wea. Rev.*, **138**, 1567-
609 1586.

- 610 Buehner, M., J. Morneau, and C. Charette, 2013: Four-dimensional ensemble-variational data
611 assimilation for global deterministic weather prediction. *Nonlin. Processes Geophys.*, **20**,
612 669-682, doi:10.5194/npg-20-669-2013, 2013.
- 613 Clayton A. M, A. C. Lorenc and D. M. Barker, 2013: Operational implementation of a hybrid
614 ensemble/4D-Var global data assimilation system at the Met Office. *Quart. J. Roy. Meteor.*
615 *Soc.*, **139**, 1445-1461.
- 616 Courtier, P., J.-N. Thépaut, and A. Hollingsworth, 1994: A strategy for operational
617 implementation of 4D-Var, using an incremental approach. *Quart. J. Roy. Meteor. Soc.*,
618 **120**, 1367-1388.
- 619 Dee, D. P., 2005: Bias and data assimilation, *Quart. J. Roy. Meteor. Soc.* **613**, 3323-3343.
- 620 Derber, J. C., and W.-S. Wu, 1998: The use of TOVS cloud-cleared radiances in the NCEP SSI
621 analysis system. *Mon. Wea. Rev.*, **126**, 2287–2299.
- 622 Desroziers, G., J.-T. Camino, and L. Berre, 2014: 4D-EnVar: Link with 4D state formulation of
623 variational assimilation and different possible implementations. *Quart. J. Roy. Meteor. Soc.*,
624 in press. doi: 10.1002/qj.2325
- 625 Errico, R.M., R. Yang, N. Privé, K.-S. Tai, R. Todling, M.E. Sienkiewicz, and J. Guo, 2012:
626 Development and validation of observing system simulation experiments at NASA's Global
627 Modeling and Assimilation Office. *Quart. J. Roy. Meteor. Soc.*, **139**, 1162-1178.
- 628 Fairbairn, D., S. R. Pring, A. C. Lorenc, and I. Roulstone, 2014: A comparison of 4DVar with
629 ensemble data assimilation methods. *Quart. J. Roy. Meteor. Soc.*, **140**, 281-294.
- 630 Gauthier, P., and J.-N. Thépaut, 2001: Impact of the digital filter as a weak constraint in the
631 preoperational 4DVar assimilation system of Météo-France. *Mon. Wea. Rev.*, **128**, 2905-
632 2919.

- 633 Gauthier, P., M. Tanguay, S. Laroche, and S. Pellerin, 2007: Extension of 3DVar to 4DVar:
634 Implementation of 4DVar at the Meteorological Service of Canada. *Mon. Wea. Rev.*, **135**,
635 2339–2364.
- 636 Gustafsson, N., 1993: Use of a digital filter as weak constraint in variational data assimilation.
637 *Proc. Workshop on Variational Assimilation, with Special Emphasis on Three-Dimensional*
638 *Aspects*, Reading, United Kingdom, ECMWF, 327–338.
- 639 Hunt, B., E. Kalnay, E. J. Kostelich, E. Ott, D. J. Patil, T. Sauer, I. Szunyogh, J. A. Yorke, and
640 A. V. Zimin, 2004: Four-dimensional ensemble Kalman filtering. *Tellus*, **56A**, 273-277.
- 641 Hunt, B. R., E. Kostelich, and I. Szunyogh 2007: Efficient Data Assimilation for Spatiotemporal
642 Chaos: a Local Ensemble Transform Kalman Filter, *Physica D*, **230**, 112-126.
- 643 Kleist, D. T., D. F. Parrish, J. C. Derber, R. Treadon, R. M. Errico, and R. Yang, 2009:
644 Improving incremental balance in the GSI 3DVar analysis system. *Mon. Wea. Rev.*, **137**,
645 1046-1060.
- 646 Kleist, D. T., 2012: An evaluation of hybrid variational-ensemble data assimilation for the NCEP
647 GFS. Ph.D thesis. Dept. of Atmospheric and Oceanic Science. University of Maryland-
648 College Park. 149 pp. (Available online at <http://drum.lib.umd.edu/handle/1903/13135>).
- 649 Kleist, D.T., and K. Ide, 2014: An OSSE-based evaluation of hybrid variational-ensemble data
650 assimilation for the NCEP GFS, Part I: System description and 3D-Hybrid results. *Mon.*
651 *Wea. Rev.*, under review.
- 652 Kuhl, D. D., T. E. Rosmond, C. H. Bishop, J. McClay, and N. L. Baker, 2013: Comparison of
653 hybrid ensemble/4DVar and 4DVar within the NAVDAS-AR data assimilation framework.
654 *Mon. Wea. Rev.*, **141**, 2740-2758.

- 655 Lawless, A. S., 2010: A note on the analysis error associated with 3D-FGAT. *Quart. J. Roy.*
656 *Meteor. Soc.*, **136**: 1094–1098.
- 657 Lewis, J. M. and J. C. Derber, 1985: The use of adjoints equations to solve a variational
658 adjustment problem with advective constraints. *Tellus*. **37A**, 309–322.
- 659 Liu C., Q. Xiao, and B. Wang, 2008: An ensemble-based four dimensional variational data
660 assimilation scheme. Part I: technique formulation and preliminary test. *Mon. Wea. Rev.*,
661 **136**, 3363-3373.
- 662 Liu C., Q. Xiao, and B. Wang, 2009: An ensemble-based four dimensional variational data
663 assimilation scheme. Part II: Observing system simulation experiments with advanced
664 research WRF. *Mon. Wea. Rev.*, **137**, 1687-1704.
- 665 Liu C., and Q. Xiao, 2013: An ensemble-based four dimensional variational data assimilation
666 scheme. Part III: Antarctic applications with advanced research WRF using real data. *Mon.*
667 *Wea. Rev.*, **141**, 2721-2739.
- 668 Lorenc, A. C., 2003: The potential of the ensemble Kalman filter for NWP – a comparison with
669 4D-VAR. *Quart. J. Roy. Meteor. Soc.*, **129**, 3183-3203.
- 670 Lorenc, A. C., and F. Rawlins, 2005: Why does 4D-Var beat 3D-Var?. *Quart. J. Roy. Meteor.*
671 *Soc.*, **131**, 3247-3257.
- 672 Lorenc, A.C. 2013: Recommended nomenclature for EnVar data assimilation methods. *Research*
673 *Activities in Atmospheric and Oceanic Modeling*. WGNE, 2pp, URL: [http://www.wcrp-](http://www.wcrp-climate.org/WGNE/BlueBook/2013/individual-articles/01_Lorenc_Andrew_EnVar_nomenclature.pdf)
674 [climate.org/WGNE/BlueBook/2013/individual-](http://www.wcrp-climate.org/WGNE/BlueBook/2013/individual-articles/01_Lorenc_Andrew_EnVar_nomenclature.pdf)
675 [articles/01_Lorenc_Andrew_EnVar_nomenclature.pdf](http://www.wcrp-climate.org/WGNE/BlueBook/2013/individual-articles/01_Lorenc_Andrew_EnVar_nomenclature.pdf))

- 676 Lorenc, A. C., N. E. Bowler, A. M. Clayton, S. R. Pring, and D. Fairbairn, 2014: Comparison of
677 hybrid-4DEnVar and hybrid-4DVar data assimilation methods for global NWP. *Mon. Wea.*
678 *Rev.*, submitted.
- 679 Lynch, P., and X.Y. Huang, 1992: Initialization of the HIRLAM model using a digital filter.
680 *Mon. Wea. Rev.*, **120**, 1019-1034.
- 681 Polavarapu, S., M. Tanguay, L. Fillion, 2000: Four-dimensional variational data assimilation
682 with digital filter initialization. *Mon. Wea. Rev.*, **128**, 2491–2510.
- 683 Rabier, F., J.-N. Thépaut, and P. Courtier, 1998: Extended assimilation and forecast experiments
684 with a four-dimensional variational assimilation system. *Quart. J. Roy. Meteor. Soc.*,
685 **124**: 1861–1887.
- 686 Rabier, F., H. Järvinen, E. Klinker, J.-F. Mahfouf, and A. Simmons, 2000: The ECMWF
687 operational implementation of four-dimensional variational assimilation. Part I:
688 Experimental results with simplified physics. *Quart. J. Roy. Meteor. Soc.*, **126**, 1143-1170.
- 689 Rawlins, F., S. P. Ballard, K. J. Bovis, A. M. Clayton, D. Li, G. W. Inverarity, A. C. Lorenc, and
690 T. J. Payne, 2007: The Met Office global 4-Dimensional data assimilation system. *Quart. J.*
691 *Roy. Meteor. Soc.*, **133**, 347–362.
- 692 Rosmond, T., and L. Xu, 2006: Development of NAVDAS-AR: nonlinear formulation and outer
693 loop tests. *Tellus*, **58A**, 45-59.
- 694 Tian, X., Z. Xie, and A. Dai, 2008: An ensemble-based explicit four-dimensional variational
695 assimilation method. *J. Geophys. Res.*, **113**, D21124, doi: 10.1029/2008JD010358.
- 696 Tian, X., Z. Xie, and Q. Sun, 2011: A POD-based ensemble four-dimensional variational
697 assimilation method. *Tellus*, **63A**, 805-816.

- 698 Wee, T.-K., Y.-H. Kuo, 2004: Impact of a Digital Filter as a Weak Constraint in MM5 4DVar:
699 An Observing System Simulation Experiment. *Mon. Wea. Rev.*, **132**, 543–559.
- 700 Wang, X., D. Parrish, D. Kleist, J. Whitaker, 2013: GSI 3DVar-Based Ensemble–Variational
701 Hybrid Data Assimilation for NCEP Global Forecast System: Single-Resolution
702 Experiments. *Mon. Wea. Rev.*, **141**, 4098–4117.
- 703 Wang, X. and T. Lei, 2014: GSI-based four dimensional ensemble-variational (4DEnsVar) data
704 assimilation: Formulation and single resolution experiments with real data for the NCEP
705 Global Forecast System. *Mon. Wea. Rev.*, **142**, 3303-3325.
- 706 Whitaker, J. S., and T. M. Hamill, 2002: Ensemble data assimilation without perturbed
707 observations. *Mon. Wea. Rev.*, **130**, 1913-1924.
- 708 Whitaker, J. S., T. M. Hamill, X. Wei, Y. Song, and Z. Toth, 2008: Ensemble data assimilation
709 with the NCEP Global Forecast System. *Mon. Wea. Rev.*, **136**, 463-482.
- 710 Whitaker, J. S. and T. M. Hamill, 2012: Evaluating methods to account for system errors in
711 ensemble data assimilation. *Mon. Wea. Rev.*, **140**, 3078-3089.
- 712 Yang, S.-C., E. Kalnay, B. Hunt, and N.E. Bowler, 2009: Weight interpolation for efficient data
713 assimilation with the Local Ensemble Transform Kalman Filter. *Quart. J. Roy. Meteor. Soc.*,
714 **135**, 251-262.
- 715 Yang, S.-C., E. Kalnay, and B. Hunt, 2012: Handling nonlinearity in an Ensemble Kalman Filter:
716 experiments with the three variable Lorenz model. *Mon. Wea. Rev.*, **140**, 2628-2646.
- 717 Zhang, M. and F. Zhang, 2012: E4DVar: Coupling an ensemble Kalman filter with four-
718 dimensional variational data assimilation in a limited-area weather prediction model. *Mon.*
719 *Wea. Rev.*, **140**, 587–600.

720 Zhu, Y., J. Derber, A. Collard, D. Dee, R. Treadon, G. Gayno, and J. A. Jung, 2014: Enhanced
721 radiance bias correction in the National Centers for Environmental Prediction's Gridpoint
722 Statistical Interpolation data assimilation system. *Quart. J. Roy. Meteor. Soc.*, **140**, 1479-
723 1492.

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<i>Experiment</i>	<i>Description</i>	<i>Relevant Equations</i>
4DVar	4DVar	Eqs. (1)-(2)
H-4DVar	Hybrid ensemble-4DVar (with adjoint) $\beta_f^{-1}=0.25, \beta_e^{-1}=0.75$	Eqs. (2)-(4)
4DEnVar	4DEnVar(no static B) $\beta_f^{-1}=0.0, \beta_e^{-1}=1.0$	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$
H-4DEnVar	Hybrid 4D-Ensemble-Var $\beta_f^{-1}=0.25, \beta_e^{-1}=0.75$	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$

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Table 1: Description of various 4D (and hybrid) single observation experiments and relevant equations for Figs. 1 and 2.

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<i>Experiment</i>	<i>Total Tendency</i>	<i>Gravity Mode Tendency</i>	<i>Ratio</i>
4DEnVar	2.853×10^{-3}	2.742×10^{-3}	0.96
4DEnVar+TLNMC	3.333×10^{-4}	2.076×10^{-4}	0.62
4DEnVar+JCDFI	1.841×10^{-3}	1.797×10^{-3}	0.98
4DEnVar+COMB	1.301×10^{-4}	9.046×10^{-5}	0.70
H-4DEnVar	2.509×10^{-3}	2.329×10^{-3}	0.93

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Table 2: The root mean square sum of the incremental (spectral) tendencies (total and gravity mode) as well as the ratio (gravity mode / total) for the eight vertical modes kept as part of the TLNMC for the single analysis valid at 06 UTC 15 July 2010 using various 4DEnVar configurations.

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<i>Experiment</i>	<i>Control</i>	<i>Description</i>	<i>Relevant Equations</i>
3DVar	--	3DVar control	Part I
3DHYB	3DVar	3D-Hybrid $\beta_f^{-1}=0.25, \beta_e^{-1}=0.75$	Part I
4DEnVar	--	4DEnVar No Constraints $\beta_f^{-1}=0., \beta_e^{-1}=1.0$	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$
H-4DEnVar	4DEnVar	Hybrid 4DEnVar No constraints $\beta_f^{-1}=0.25, \beta_e^{-1}=0.75$	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$
H-4DEnVar_NMI	H-4DEnVar 3DHYB	H-4DEnVar + TLNMC (all observation k -levels)	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$ Eq. (7)
H-4DEnVar_DFI	H-4DEnVar	H-4DEnVar + JCDFI ($\chi=10$)	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$ Eqs. (8)-(9)
H- 4DEnVar_COMB	H-4DEnVar	H-4DEnVar + TLNMC (k =mid only) + JCDFI ($\chi=10$)	Eqs. (3) & (5) $\mathbf{M}_k=\mathbf{I}$ Eq. (7) for $k=m$ Eqs. (8)-(9)

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Table 3: Description of various (hybrid) OSSE-based experiments.

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760

761 **List of Figures**

762 **Figure 1.** 500 hPa Temperature (shaded, K) and vector wind (vectors, m s^{-1}) analysis increment
763 resulting from the assimilation of a single 500 hPa temperature observation (-2K innovation, 1K
764 error) taken -3h from analysis time (location of observation denoted with black dot) for 4DVar
765 (upper left), H-4DVar (lower left), 4DEnVar (upper right), and H-4DEnVar (lower right). The
766 solid contours are the background 500 hPa geopotential height (dm) valid at the center of the
767 assimilation window (analysis time). The reference arrow is representative of 1 m s^{-1} .

768
769 **Figure 2.** As in Fig. 1, except for the analysis increment valid at the beginning (t-3h, top), center
770 (t=0h, middle), and end (t+3h, bottom) of the assimilation window for the H-4DVar (left) and H-
771 4DEnVar (right) cases. The solid contours are the background 500 hPa geopotential height (dm)
772 valid at beginning (top), center (middle), and end (bottom) of the assimilation window.

773
774 **Figure 3.** Power spectra of the divergence increment for a single analysis valid at 06 UTC 15
775 July 2010 on model level 35 (top, approximately 200 hPa) and 14 (bottom, approximately 850
776 hPa) utilizing 4DEnVar (non-hybrid) with no constraints (red), tangent linear normal mode
777 constraint (blue), weak constraint digital filter (purple), and combined normal mode and weak
778 constraint (green).

779
780 **Figure 4.** Observation penalty (top) and reduction of the gradient norm (bottom) by iteration for
781 an analysis valid at 06 UTC 15 July 2010 utilizing 4DEnVar (non-hybrid) with no constraints
782 (red), tangent linear normal mode constraint (blue), weak constraint digital filter (purple), and
783 combined normal mode and weak constraint (green).

784 **Figure 5.** Difference in the time-averaged zonal mean standard deviation of the analysis error
785 for the H-4DEnVar_NMI experiment relative to 3DHBYB for zonal wind (top, ms^{-1}), temperature
786 (middle, K), and specific humidity (bottom, g kg^{-1}), covering the period spanning 00 UTC 1
787 August 2005 through 18 UTC 30 August 2005.

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789 **Figure 6.** Difference in the total energy (J kg^{-1}) time-averaged global mean standard deviation
790 of the analysis error for the H-4DEnVar_NMI (blue), H-4DEnVar_DFI (purple), and H-
791 4DEnVar_COMB (green) experiments relative to H-4DEnVar, covering the period spanning 00
792 UTC 1 August 2005 through 18 UTC 30 August 2005.

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794 **Figure 7.** Time series of the global mean surface pressure error (hPa) for the background (top)
795 and analysis (bottom) from the H-4DEnVar (red), H-4DEnVar_NMI (blue) and H-
796 4DEnVar_DFI (purple) experiments.

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798 **Figure 8.** The root mean square sum of the incremental spectral tendencies (top) for the total
799 tendency (dashed) and gravity mode tendency for the eight vertical modes (solid) for the
800 4DEnVar (red), 4DEnVar_NMI (blue), 4DEnVar_DFI (purple) and 4DEnVar_COMB (green)
801 configurations for the single analysis valid at 06 UTC 15 July 2010 as a function of observation
802 bin (analysis relative time). Also plotted is the ratio (bottom, gravity mode/total tendency) for
803 the 4DEnVar (red), 4DEnVar_NMI (blue), 4DEnVar_DFI (purple) and 4DEnVar_COMB
804 (green) configurations.

805

806 **Figure 9.** As in Fig. 5, but for the difference H-4DEnVar-4DEnVar.

807 **Figure 10.** Time-averaged 500 hPa anomaly correlation (upper panels) for the Northern
808 Hemisphere (left) and Southern Hemisphere (right) for the 3DHYB (red), H-4DEnVar_NMI
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810 lead time as well as the difference (lower panels) for H-4DEnVar-3DHYB (blue) and 3DVar-
811 3DHYB (green). The 95% confidence threshold for a significance test (derived from a standard
812 t-test) is also plotted in the lower panels for the H-4DEnVar (blue bars) and 3DVar (green bars)
813 differences.

814
815 **Figure 11.** Time-averaged root mean square geopotential height errors (m) for forecasts from
816 the 00 UTC analyses in the 3DHYB experiment as a function of lead time for the Northern
817 Hemisphere (upper left) and Southern Hemisphere (lower left) verified against the ECMWF
818 nature run forecasts verifying between 27 July 2005 and 01 September 2005. The difference
819 between the two experiments (H-4DEnVar_NMI-3DHYB) for the Northern Hemisphere (upper
820 right) and Southern Hemisphere (lower right) is also plotted.

821
822 **Figure 12.** Percent change in root mean square error from H-4DENVAR_NMI minus 3DHYB
823 for the period covering 27 July 2005 and 01 September 2005 in the northern hemisphere (green),
824 tropics (red) and southern hemisphere (blue) for selected variables as a function of forecast lead
825 time. The forecast variables include 500 hPa geopotential height (upper left, upper right), 850
826 hPa vector wind (upper middle), 200 hPa vector wind (middle row), and 700 hPa specific
827 humidity (bottom row). All verification is performed against the ECMWF nature run. The error
828 bars represent the 95% confidence threshold for a significance test.

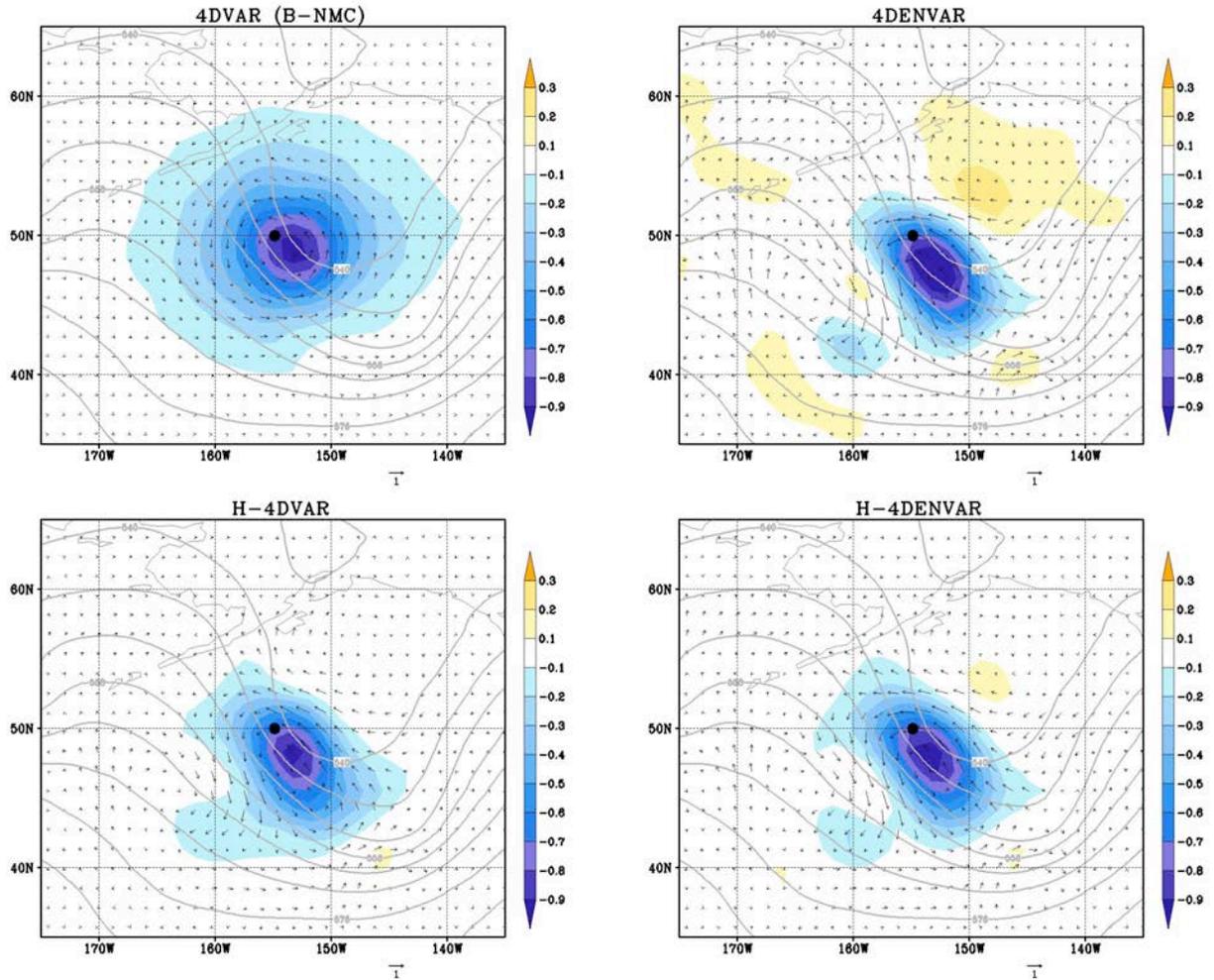


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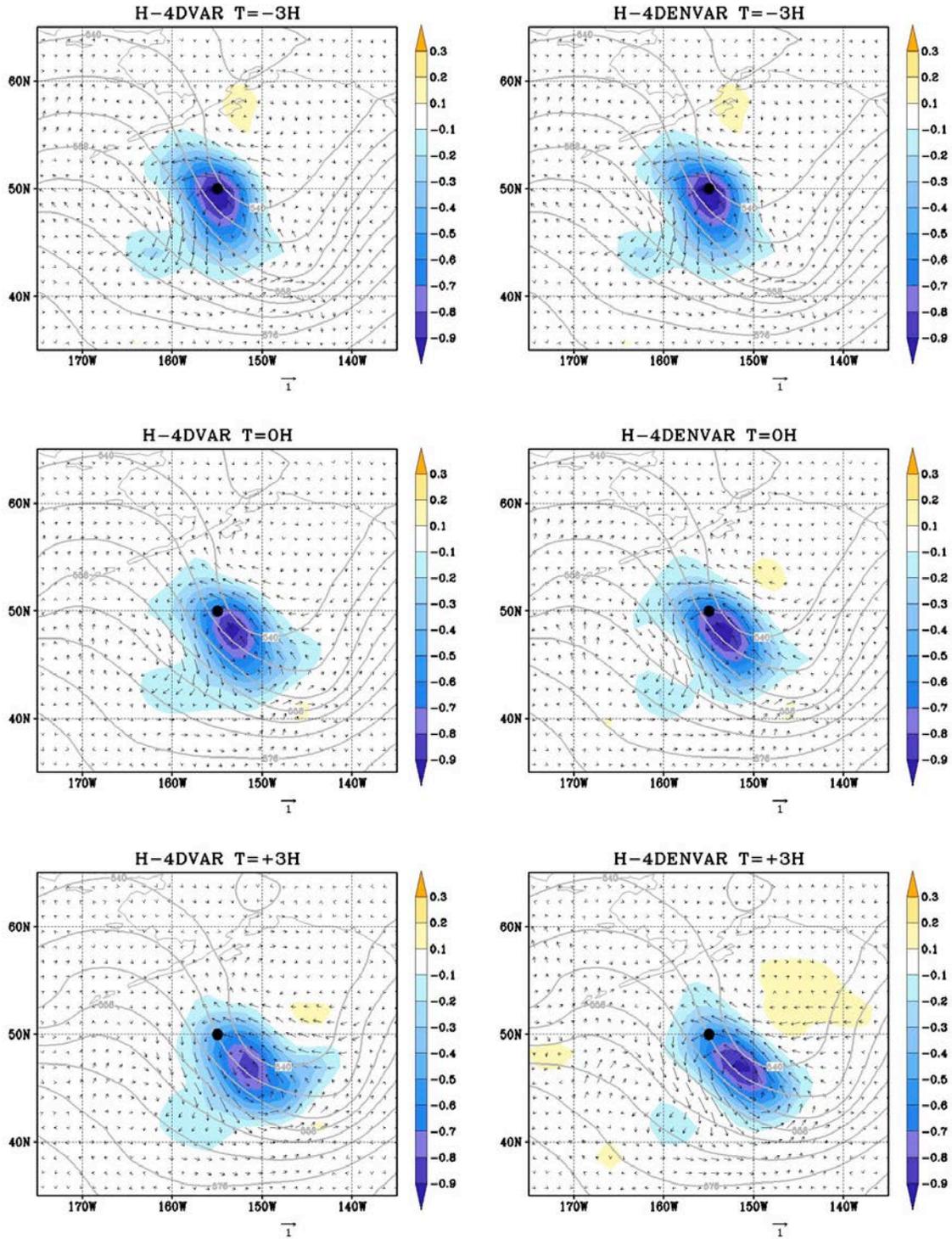


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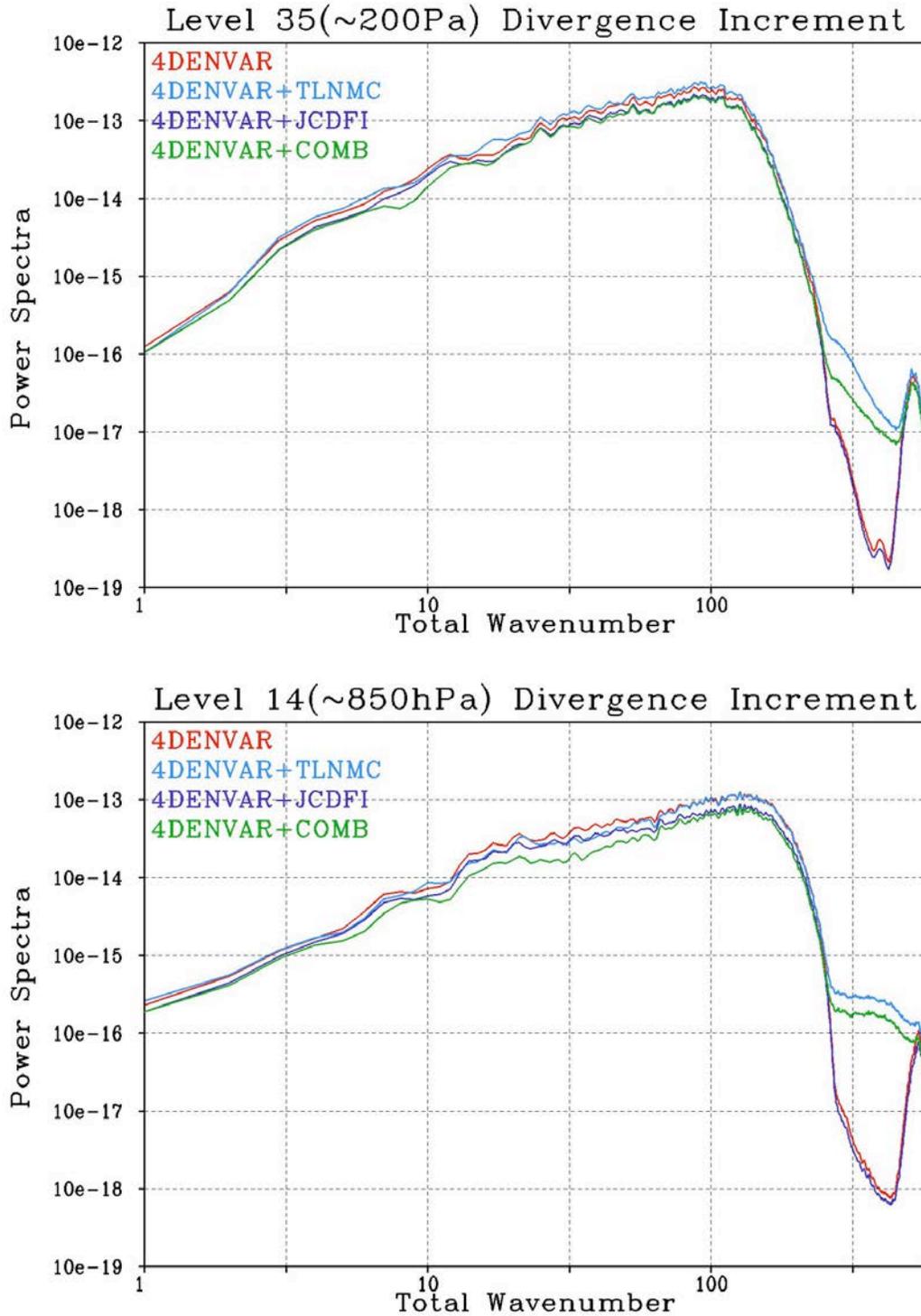


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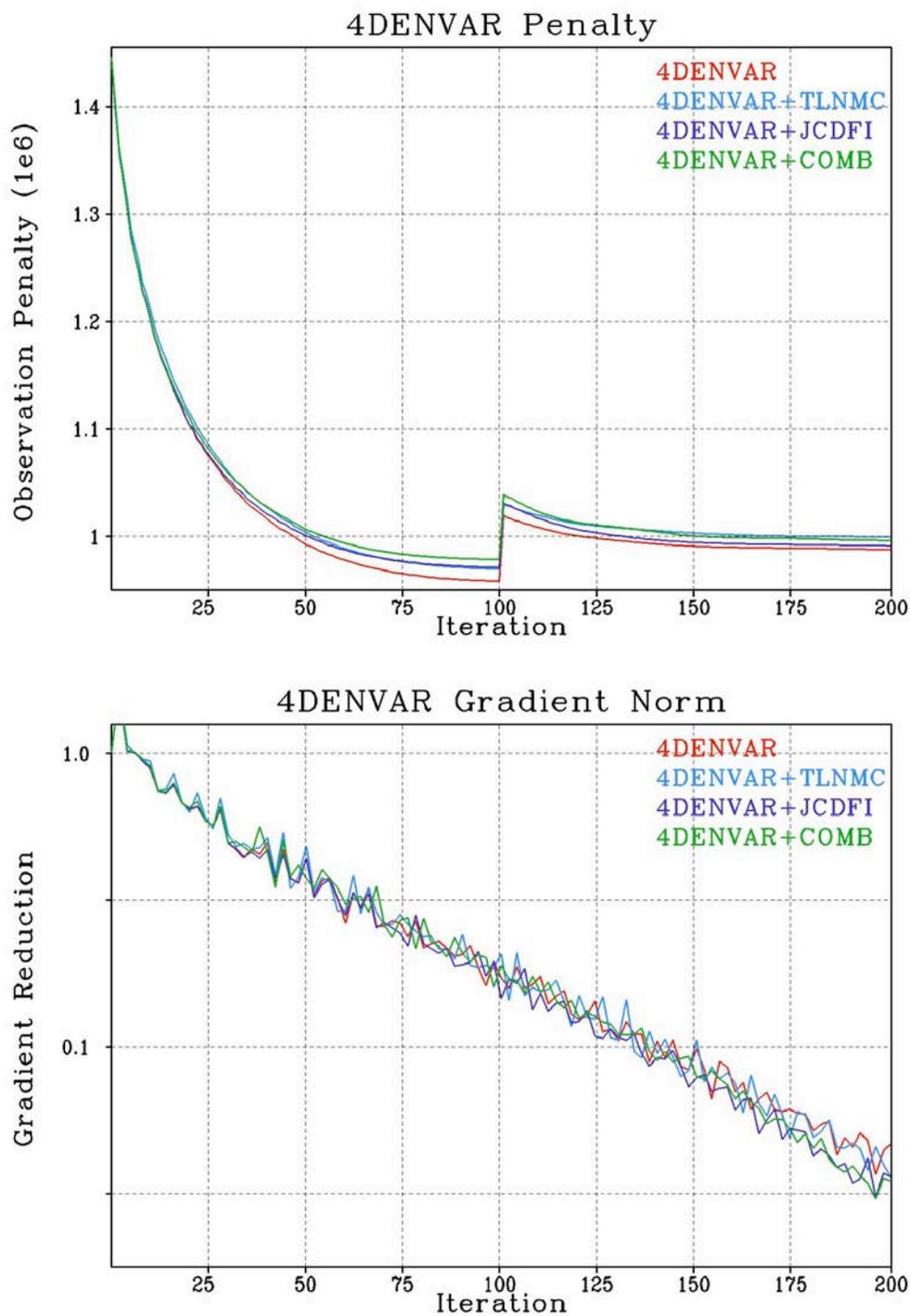


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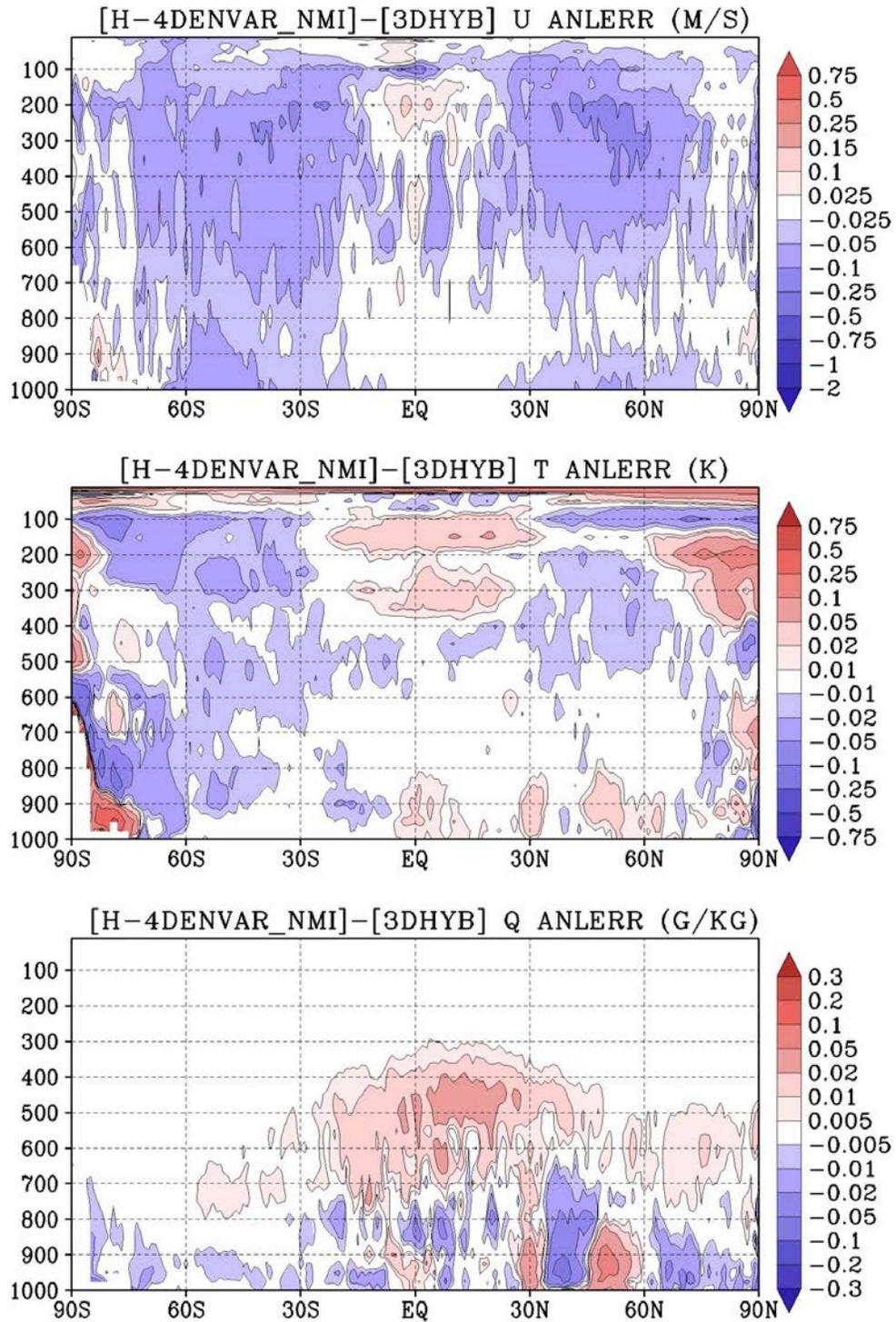


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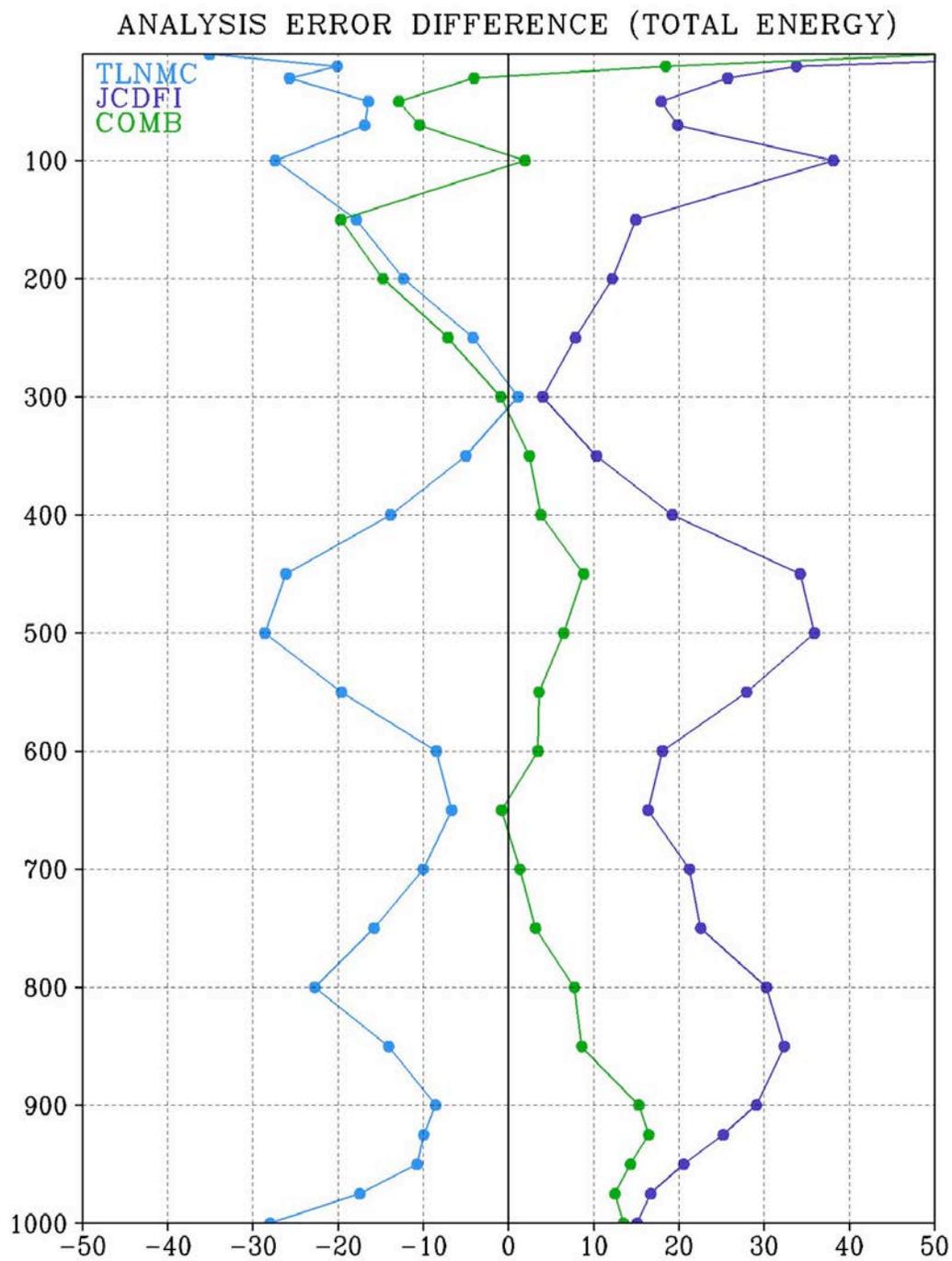


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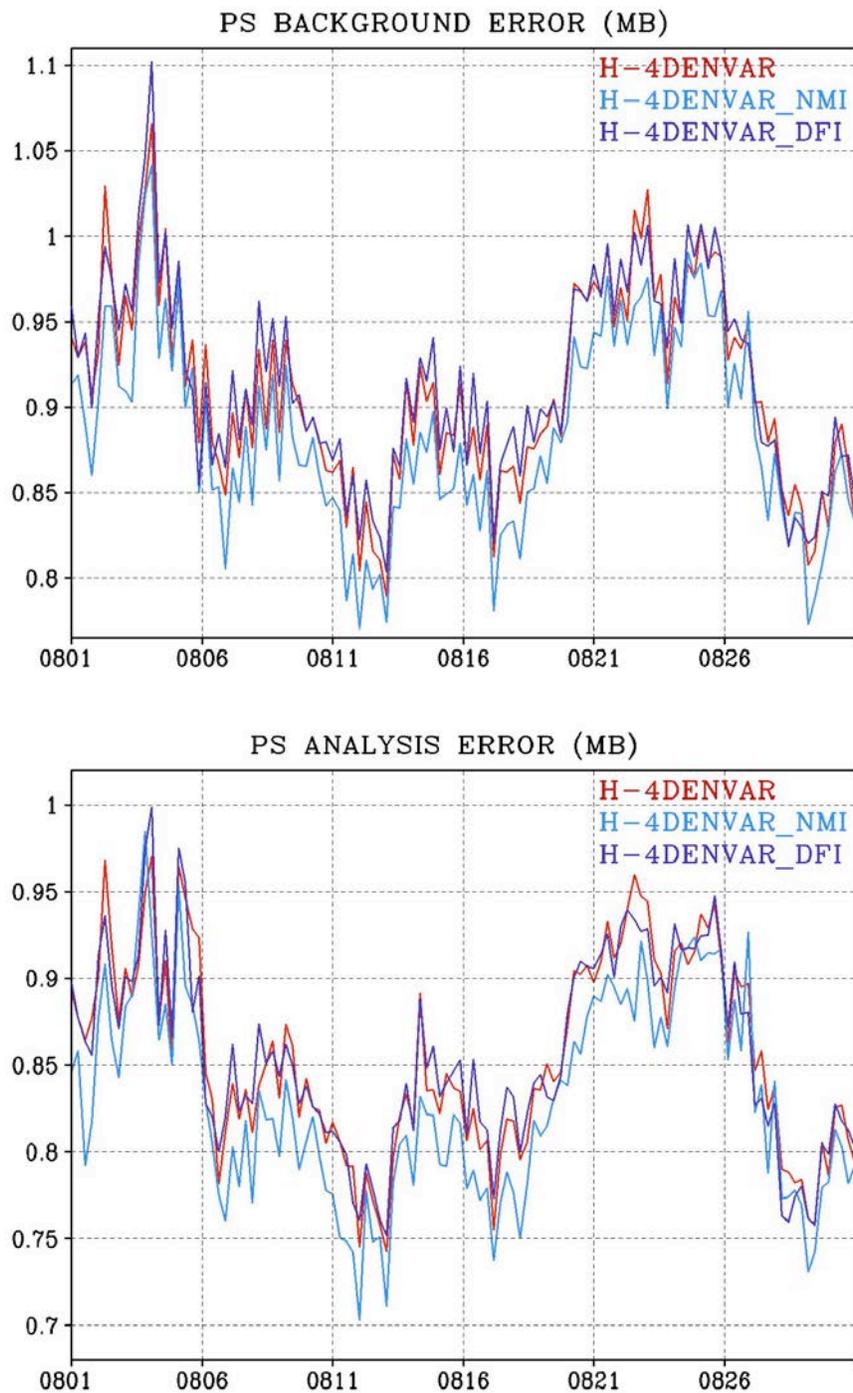


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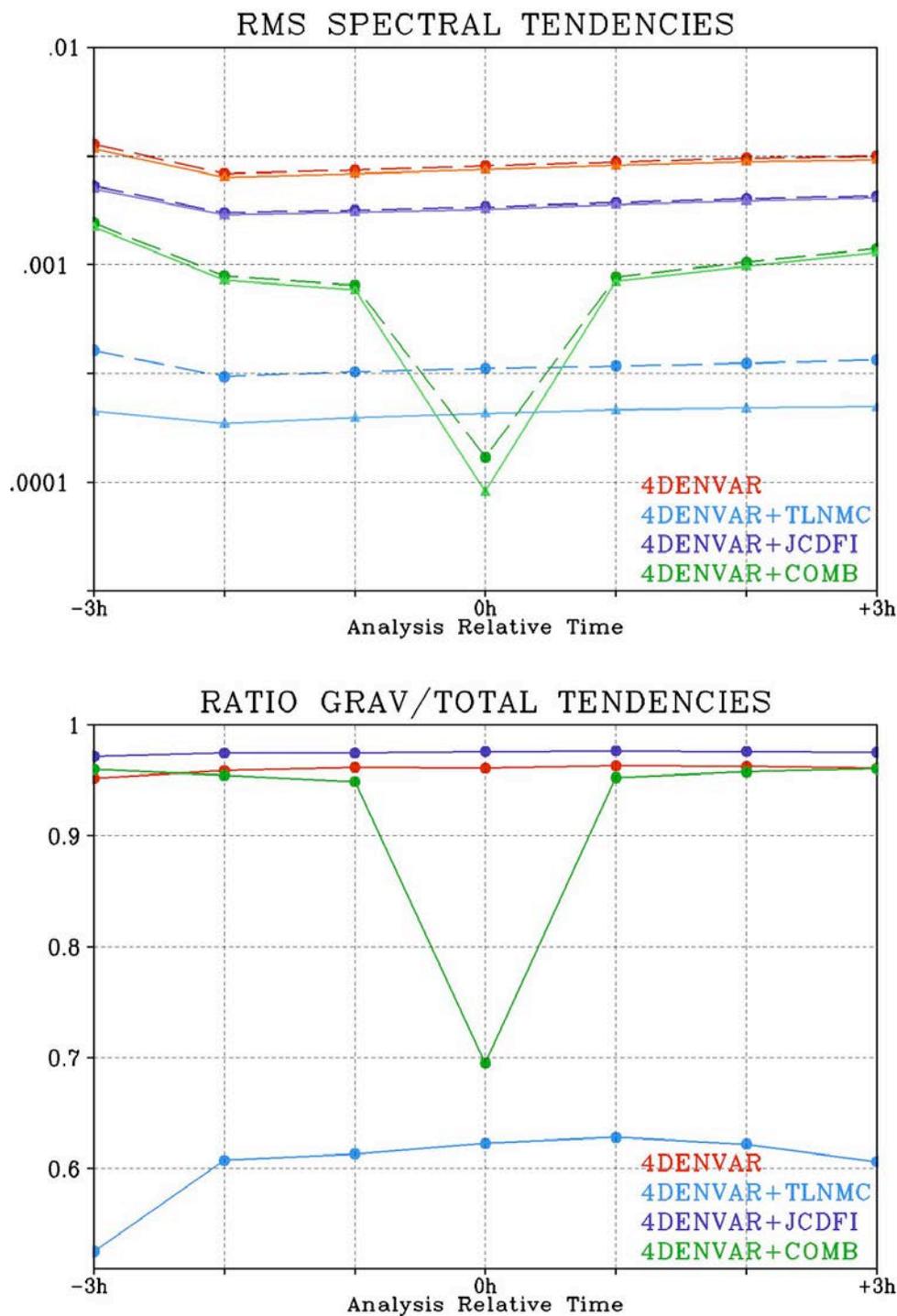


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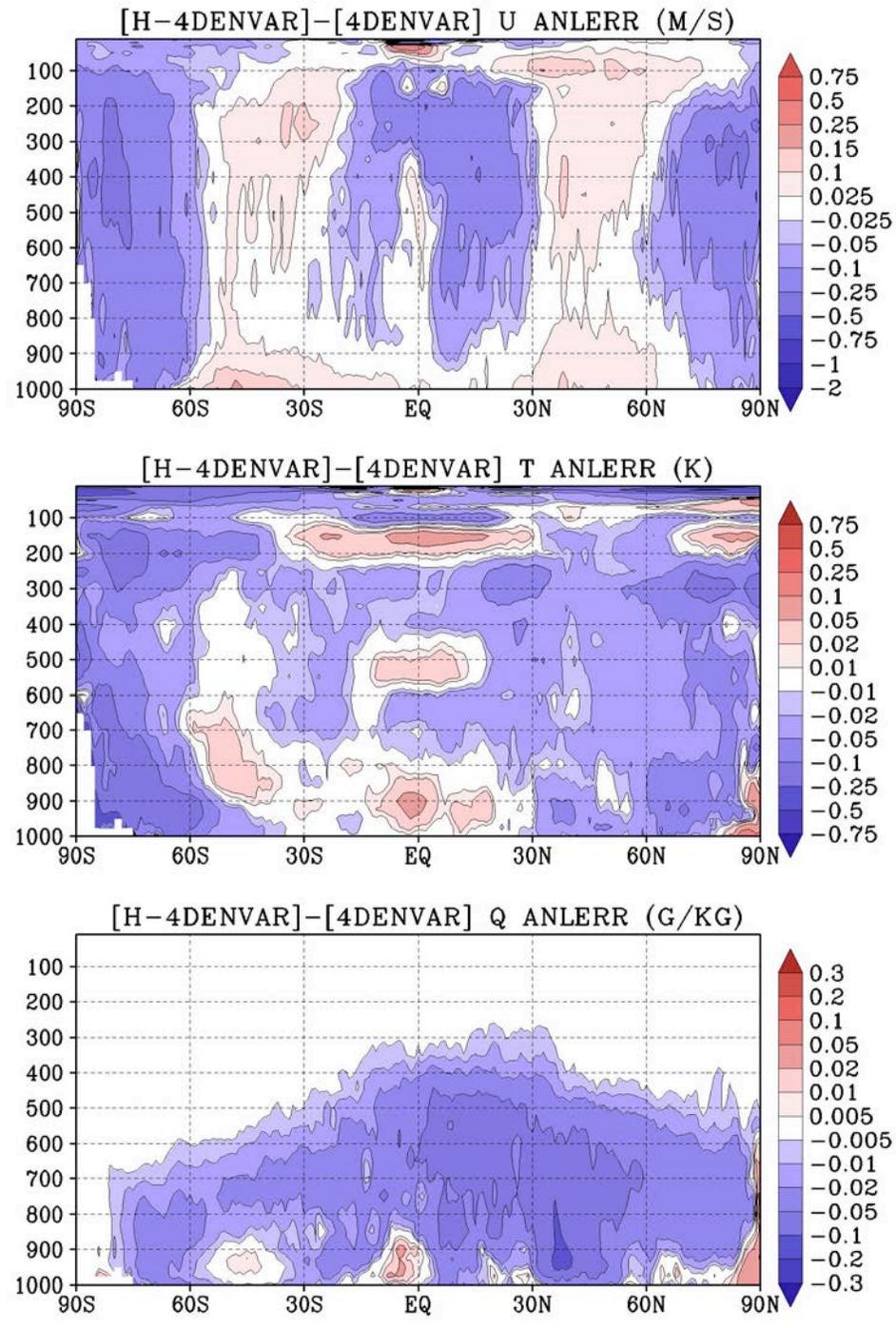


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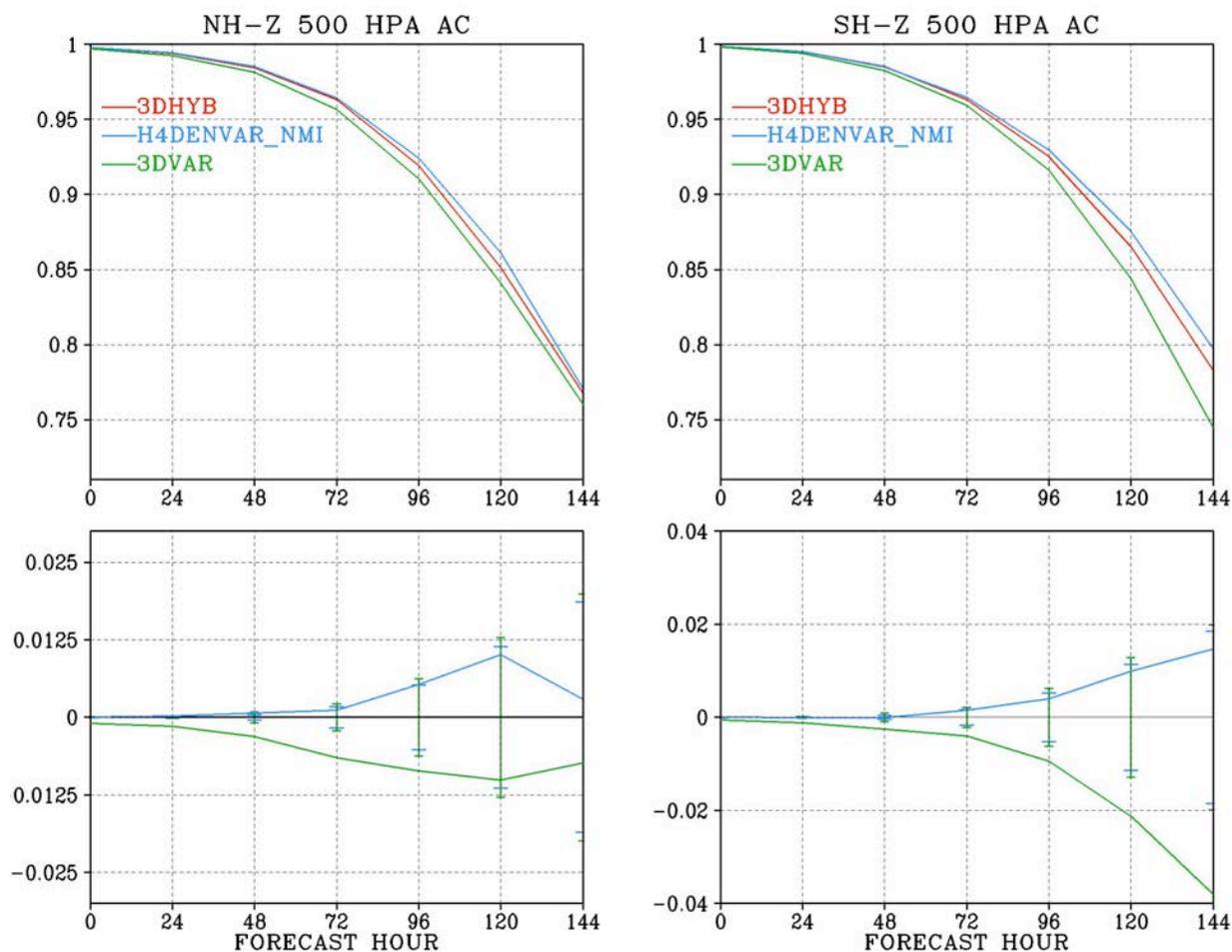


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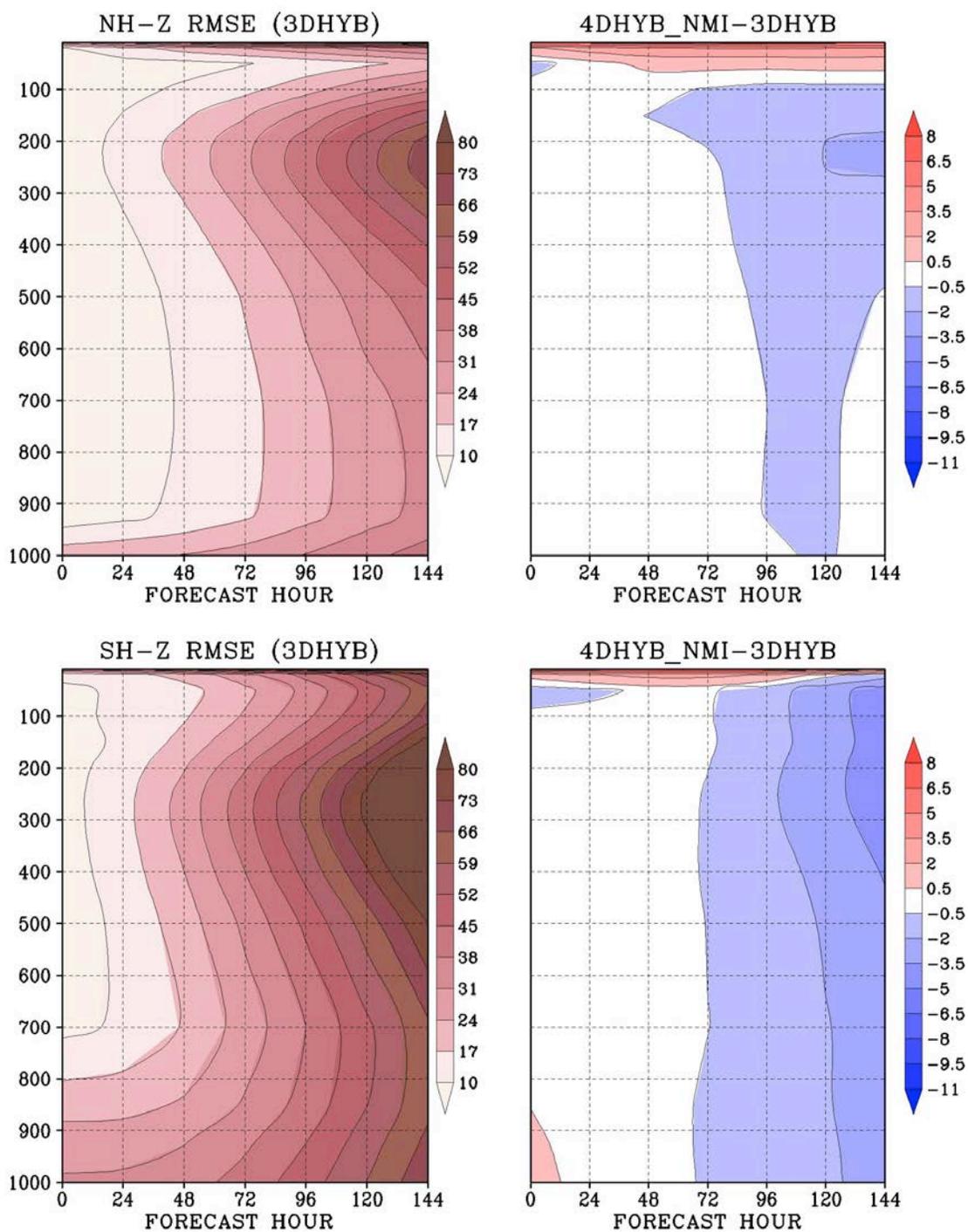


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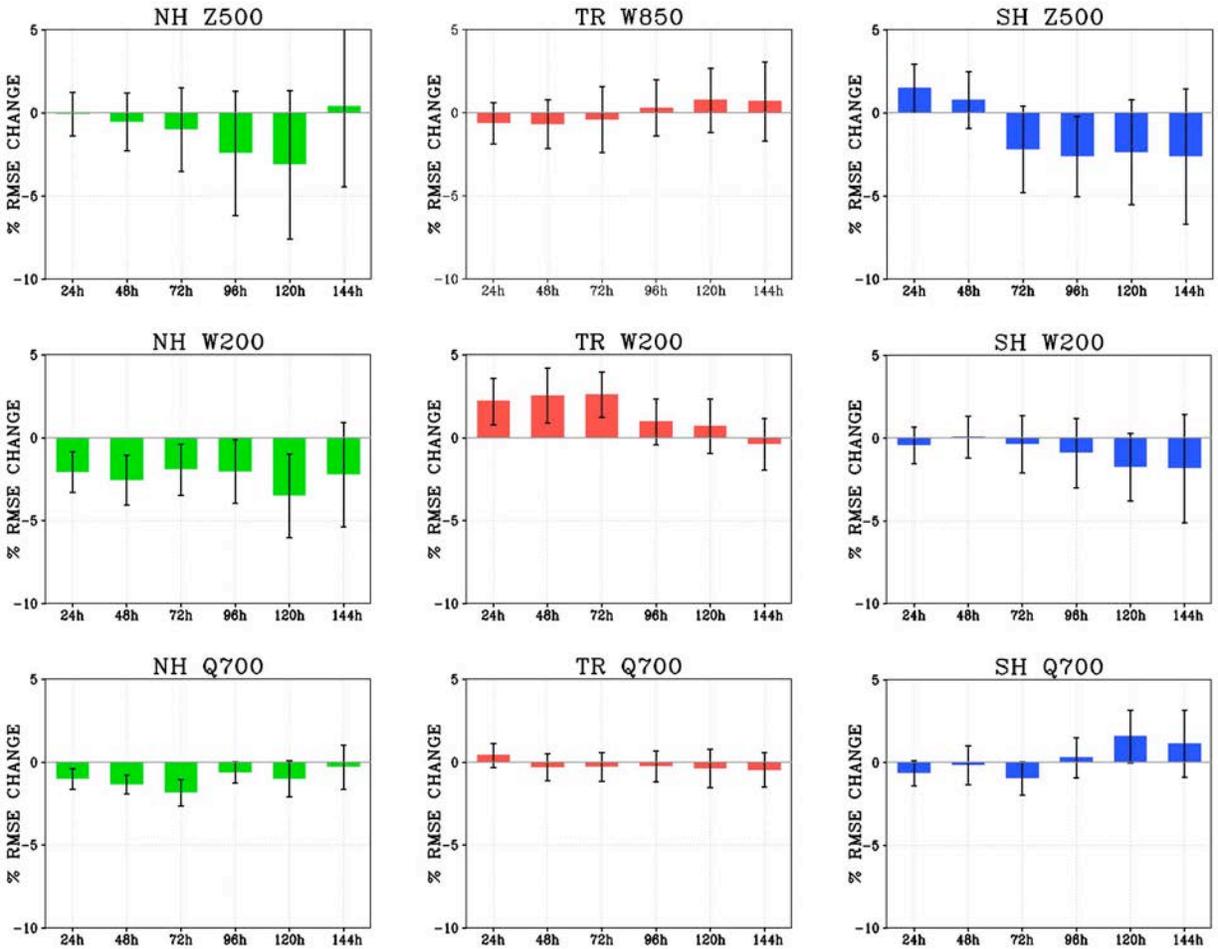


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